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Gaussian distributions on a logarithmic scale. As in England, width commonly exceeds length. Vertical dimensions correlate with length more than with width. Cirque form varies with geology, but also with relief as both vary between mountain groups. The main contrast is between larger, better-developed cirques and higher relief on volcanic rocks in the north-west, and smaller, less-developed cirques and lower relief on sedimentary rocks in the south.

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**Allometric development of glacial cirque form:  
geological, relief and regional effects on the cirques of Wales**

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**Abstract**

Headward and downward erosion near glacier sources, at rates exceeding fluvial erosion, is important in recent discussions of orogen development and the limits to relief. This relates to a long history of debate on how the form of glacial cirques develops, which can be advanced by relating shape to size in large data sets. For 260 cirques in Wales, this confirms different rates of enlargement in the three dimensions: faster in length than in width, and slower in vertical dimension whether expressed as overall height range, axial height range or wall height. Maximum gradient, plan closure and number of cols increase with overall size. This allometric development applies over different cirque types, regions and rock types. Headwall retreat, often by collapse following glacial erosion at the base, is faster than downward erosion. Welsh cirques form a scale-specific population and, as in other regions, size variables follow Gaussian distributions on a logarithmic scale. As in England, width commonly exceeds length. Vertical dimensions correlate with length more than with width. Cirque form varies with geology, but also with relief as both vary between mountain groups. The main contrast is between larger, better-developed cirques and higher relief on volcanic rocks in the north-west, and smaller, less-developed cirques and lower relief on sedimentary rocks in the south.

**Keywords:** Glacial erosion; Cirque form; Allometry; Morphometry; Wales; Statistical graphics.

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## 1. Introduction

### 1.1. Aims and objectives

The erosion of glacial cirques takes tens to hundreds of thousands of years and is thus not susceptible to direct observation. Clearly cirques start from glacial occupation of already indented topography, often fluvial but sometimes volcanic or following landsliding. Several schemes have been proposed for cirque development (e.g. Gordon, 1977; Bennett, 1990): all are based on the concept that cirques enlarge over time, as the bedrock removed cannot be replaced in the same form. This substitution of spatial variation for change over time is known as the ergodic hypothesis (Cox, 1977; Thorn, 1988).

Cirque distribution is excellent evidence of former glaciation and is fairly well established, but data on the precise form of glacial cirques are limited and based on varied definitions. Here I report on an effort to provide comprehensive and comparable data, relevant to hypotheses of cirque development.

The aim of this paper is to use extensive data on spatial variation in cirque form to show how cirques may develop. This involves relating both morphometric and qualitative data to a number of possible controls: geology, relief, position and timing or type of last glacial occupation. The objectives are to present a high-resolution data set for all cirques in Wales, to use it to test hypotheses of changes in cirque shape with size and of environmental controls on both, and to set the statistical results in the context of a long historical debate on downward and headward cirque erosion. The data set is sufficiently large to be subdivided on various criteria, with numerous cirques in each class of geology or grade of development. This permits more reliable models, narrower confidence intervals, and statistical tests of greater power. The data also provide a sampling frame for future detailed studies of glacial morphology and Quaternary chronology, comparable to that provided for the English Lake District by Evans and Cox (1995).

### 1.2. Review of cirque development models: wearing down and wearing back

There is no obvious reason why downward erosion by glacial plucking, abrasion and meltwater should be equal to headwall recession by these and other processes (frost action and slope collapse). Contradictory views have been published, and it is worth considering the development of ideas on this controversial subject before presenting morphometric data permitting a test of their applicability. Fuller historical depth is provided in Evans (in press).

In the early twentieth century, cirques were considered to wear back by 'sapping' of the base of the headwall, usually by frost weathering and rockfall. Local glaciers flatten and deepen cirque floors, but by removing rockfall they constantly attack headwalls at their bases (de

Martonne, 1901). Richter (1900) emphasized cirque enlargement by headwall retreat through frost weathering, with glaciers abrading cirque floors but protecting them from fluvial incision; this produced a high terrace which could eventually truncate a mountain range. Richter's hypothesis that in Norway this retreat had gone so far as to create the high mountain plateaus (the Palaeic surface) was regarded as inconceivable by Ahlmann (1919, p. 220).

Johnson's (1904) bergschrund hypothesis explained headwall retreat by vigorous frost weathering in a narrow zone. This provided debris that aided glacial abrasion downstream. Hobbs (1910, 1911) proposed four stages of sculpture by mountain glaciers, from youthful 'channelled and grooved' uplands through adolescent 'early fretted' and mature 'fretted' to senile 'monumented' upland, as cirques enlarged laterally at the expense of 'pre-glacial' upland. This 'Cycle of Mountain Glaciation', emphasizing headward recession, was discussed at some length in Embleton and King (1968). It involves, however, a dangerous substitution of time by space, the more so as Hobbs' main examples of the four stages are widely separated over space, respectively in the Bighorn Mountains of Wyoming, in northwest Wales, in the (Swiss) Alps, and in Glacier National Park, Montana. Hobbs suggested that cirques become more complicated with age: others attribute the complexity of broader cirques to coalescence. Davis (1911, pp. 56-61) described the enlargement and coalescence of cirques in Colorado and supported Richter's idea of mountain truncation at the snowline by prolonged glaciation.

Blache (1952, pp. 112-5) rejected Hobbs' four stages: instead of duration of glaciation, he related such differences to increased altitude, giving greater slope and greater ice discharge. The stages seem to relate also to dissection of the initial topography. Blache (1960, p. 30) denied that cirques expand laterally as in Hobbs' scheme. Derbyshire and Evans (1976) related Hobbs' 'stages' to tectonic environment and degree of dissection, especially the spacing of valleys.

Lewis (1938, 1940) suggested that meltwater forced its way down far below the bergschrund, explaining the great height of some cirque headwalls (over 300 m, even in Britain). Following later fieldwork, in the Jötunheimen, Norway (Clark, 1951; Clark and Lewis, 1951), Lewis emphasized the importance of rotational flow in cirque development (Lewis, 1949, 1960). Rotational flow maximises basal sliding but requires a certain range of ice surface gradients (see section 3.2). Lewis (1949) calculated that gradients of 15 - 20° and ice thicknesses of 80 - 100 m gave the greatest probabilities of rotational slip, and thus erosion of rock basins, both in cirques and below valley steps. Rotational flow has now been clearly demonstrated in Norway on Veslgjuv-breen (Grove, 1960) and Vesl-Skaubreen (McCall, 1960), and in Colorado on the Arapaho Glacier (Waldrop, 1964). Overall centre-line gradients were some 17°, 27° and 19° respectively. Lewis (1960) calculated that the 63 m deep rock basin of Blea Water, in a cirque in the English Lake

District, was excavated by a rotationally flowing glacier sloping at  $13^{\circ}$  to  $16^{\circ}$ , and that at least  $7^{\circ}$  was required to overcome basal friction.

A few authors have given more prominence to the downward erosion of cirque floors. Strøm (1945) contrasted the rapid cirque backwall recession in the quartzite of Rondane (southeast Norway) with the apparent vertical erosion (producing rock basins up to 170 m deep) in the plutonic rocks of Moskenesøy, Lofoten Islands. Galibert (1962, p. 16) proposed that in the Alps, crest retreat could be negligible compared with vertical incision of glacial cirques where jointing was most pronounced. Incision was due to greater abrasion under thick ice, accompanied by collapse of the headwall base by pressure release on deglaciation (see also Birot, 1968): high summits are deeply frozen and ‘paralysed’.

Referring mainly to the Uinta Mountains of Utah, and to some Norwegian cirques, White (1970, p. 123) maintained that “In many mountainous areas... arêtes between opposing cirques are so narrow and steep that very slight additional headward erosion by either of the opposing glaciers would have breached the arêtes to form a col or a pass. Yet these arêtes persist, in rarely broken continuity threading their precarious way for great distances between the headwalls of successions of opposing cirques”. Avalanche ravines on some headwalls would not have survived undermining by active sapping. White concluded that cirque floors are lowered by glacial erosion more rapidly than walls are cut back, a view that provides the opposite extreme to Richter.

Evidence of strong headward erosion was found in south-central Alaska by Tuck (1935), who proposed divide migration of up to 1 km due to stronger glacial erosion in northward-facing valley heads; the effects of differential insolation on glacier balance were reinforced by south winds. A similar interpretation of southward divide migration had been applied in west-central British Columbia where north-facing glaciers “have cut short northerly-trending valleys on the north side of the higher mountains” (Hanson 1924, p.31). Evans (1972) confirmed Hanson’s hypothesis for several ranges in the Coast Mountains of southwest British Columbia. The main divides of the Bendor and Tatlow Ranges appear to have been shifted 1 to 2 km southward by cirque headwall retreat.

Evans (1997) emphasized the role of slope collapse in cirque headwall retreat, and listed examples of historical rock avalanches from cirque headwalls. The 1873, 1991 and other collapses around Mount Cook can be added to this list. These were controlled by stress-release joints slightly steeper than the cliffs, in cohesionless, closely jointed rock. “Many occur high on mountain slopes where gravitational collapse is the only operative erosion process to keep pace with glacial and fluvial valley incision” (McSaveney, 2002, p. 69). In the Ben Ohau Range, also in New Zealand, Brook et al. (2006) suggest that cirques lengthen and deepen faster than they widen, increasing in

maximum gradient and reaching well-developed shapes after 600 ka. Brocklehurst and Whipple (2002) suggested that cirque headwall retreat is important on the east slope of the Californian Sierra Nevada, causing divide recession of up to 2.5 km beyond that expected from fluvial erosion.

Two recent publications based on analyses of DEMs (Digital Elevation Models) emphasize the importance of glacial erosion in cirques, in mid-latitude mountains. In part of the Washington Cascades, Mitchell and Montgomery (2006) note a parallelism between three trends: non-volcano summit altitudes, cirque floors and modern glacier median altitudes. Each rises eastward at 9 to 15 m km<sup>-1</sup>. Only about 10% of each subdivision rises above the highest cirque floors, and few peaks rise more than 600 m above. Rock exhumation rates suggest vertical erosion of 2 to 5 km in the last 15 Ma, and are greatest 30 to 40 km west of the highest peaks. This is not the pattern expected from fluvial or slope erosion, so Mitchell and Montgomery propose a ‘glacial buzzsaw’ of greatest glacial erosion at the average Quaternary glacial equilibrium line represented by the cirque floors, where ice discharge and velocity were greatest. This increased the slope gradients above cirque floors to over 30°, causing slope failure. Thus both vertical and headward erosion in cirques is considered to dominate landscape development at high altitudes.

Uplift rates are greater in the Kyrgyz Range of the Tien Shan, where Oskin and Burbank (2005) use the sub-Cenozoic unconformity to suggest an east - west spatial gradient of uplift and thus of landform development. As mountains are taken above the snowline, glacial erosion both deepens and widens fluvial valleys, increasing local relief. This starts from the north slope where the snowline is some 200 m lower, pushing the divide 0.9 to 4.4 km southward. Erosion is localized at the bases of cirque headwalls, and cirque headwall retreat is two to three times the rate of vertical erosion. Glacial erosion is thus not simply a function of ice flux, but requires further help, perhaps from subglacial water pressure fluctuations in cirques. “Cirque retreat can effectively bevel across an elevated alpine plateau ...” (Oskin and Burbank, 2005, p. 936).

Discussions of cirque development often assume that the cirque is isolated. In many areas, however, cirque headwalls intersect: cirques are side by side along one side of a ridge or more rarely (where glaciation has been more symmetrical) back to back across a ridge. This may limit further enlargement of cirques despite ongoing erosion, as arêtes are lowered along with cirque floors. Also cirque thresholds may recede by erosion of a glacial trough downstream. Most authors regard both modes of development as important, and Evans (1997) emphasized the need for erosion to be greatest at the base of the headwall, to account for development of the characteristic break of slope and a low-gradient floor. This may be facilitated by water pressure fluctuations, or by a transition from cold to warm, sliding ice as the glacier thickens: both mechanisms encourage plucking of blocks of rock. Erosion of a rock basin is even stronger



evidence of glacial erosion and probably requires rotational flow, but only a minority of cirques have a rock basin or even a lake (Derbyshire and Evans, 1976). If cirque development proceeds differently in different regions, the factors controlling this need to be established. Cirques are eroded both downward and headward. They are initiated by glaciers filling hollows of diverse types, and positive feedbacks (Graf, 1976) prevent them from escaping this inheritance: they remain diverse (Evans and Cox, 1995).

[Fig. 1 about here]

## 2. Data

### 2.1. Study Area: Wales

“I cannot imagine a more instructive and interesting lesson for any one who wishes (as I did) to learn the effects produced by the passage of glaciers, than to ascend a mountain like one of those south of the upper lake of Llanberis...convex domes or bosses of naked rock, generally smoothed, but with their steep faces often deeply scored in nearly horizontal lines...” (Darwin, 1842, p. 188). Darwin’s early work on the glaciation of Wales (in the classic area of northern Snowdonia) is of historical importance in the acceptance of the glacial theory, and is set in its full context by Herbert (2005, pp. 277-284).

Choice of Wales as a study area provides a large number of cirques on a variety of rock types from Cambrian to Carboniferous (Silesian) in age, and in relief varying from the narrow mountainous ridges of Snowdon, through the more massive Carneddau Mountains, to the plateaus of central Wales and the sandstone escarpments of the south. Wales is on the west coast of Great Britain, between England and Ireland, and between 51.4 ° and 53.4 ° N. The morphology of Wales defies brief summarization, but the highest ground is in the northwest (northern Snowdonia), followed by the Aran – Cadair Idris Range in the western part of north-central Wales, and the Brecon Beacons in the south (Fig. 1). Wales and its English borderland form the Cambrian Massif, a Caledonian massif with a strong northeast-southwest grain and a Hercynian accretion in the south with a roughly east-west grain. Thomas (1970) has shown many structural influences in the topography. Quaternary tectonic movements are believed to have been subdued.

Position is used as a surrogate for climate (Evans, 1999) as there are few direct climatic observations in cirques and none for the glacial periods when they were developing. In Wales temperature varies mainly with altitude (Sumner, 1997), but at present precipitation varies with exposure, related mainly to distance from the west and south coasts but with a lag so that the wettest mountains (Snowdon, Arenigs, Arans, Pumlumon, the Rhondda area and the western Brecon

Beacons) are 20 to 30 km from the coast. In areas with cirques, modern annual precipitation (1916-50) varied from 4500 mm on Snowdon and 2500 mm on Pumlumon (also spelled Plumlumon and Plynlimon) and the western Brecon Beacons to 1300 mm in the southeast (in the Black Mountains, the Radnor Forest, and near Abergavenny), and in Yr Eifl in northwest Wales (MHLG, 1967). The driest mountains are the Clwydian Range in the northeast, which show no clear cirques. Almost all areas with modern annual precipitation over 2000 mm support glacial cirques. Precipitation in the winter half-year is mainly 54-58% of the total on the mountains of North Wales, and 57-61% on South Wales mountains.

[Fig. 2 about here]

British and Irish cirques are believed to have developed over a series of glaciations, each intermediate in extent between the present non-glacial conditions and glacial maxima when most of Wales, Ireland, Scotland and northern England was covered by a coalescent ice sheet. Wales has undergone both ice-cap and local glaciation (Campbell and Bowen, 1989; Lewis and Richards, 2005). 'Irish Sea' ice from Scotland and Cumbria covered the north coast and western peninsulas, but the local Welsh Ice Cap was strong enough to prevent exotic ice from penetrating the areas with cirques: it built up to at least 850 m altitude from Snowdon to Carnedd Llewelyn (McCarroll and Ballantyne, 2000), and 750 m around Cadair Idris (Ballantyne, 2001) and the Rhinog mountains. Jansson and Glasser (2004) recognize varying ice flow patterns, and extents of cold-based ice, during build-up, maximum, and decline of the Last Glaciation. Cold ice covered the highest summits at the maximum. Ice streams formed during deglaciation, but did not affect areas with cirques (Jansson and Glasser 2004, Fig. 3). As these areas were mainly ice sources, and often covered by ice frozen to the bed, cirques suffered little modification by ice-cap glaciation except in the Moelwyn, Rhinog and Migneint areas, between Snowdon and Cadair Idris.

Wales covers 20,760 km<sup>2</sup>, and the 260 cirques occur in an area 180 x 50 km (Fig. 2). Thus the distribution is much less dense than in the Maritime Alps, where Federici and Spagnolo (2004) measured 432 glacial cirques in 67 x 26 km (1742 km<sup>2</sup>). There is, however, a large gap between the cirques of Mid Wales and those of the Brecon Beacons, and there are few cirques in northeast Wales. The greatest concentrations are in northern Snowdonia (103 cirques in 30 x 18 km), Aran – Cadair Idris (48 cirques in 30 x 11 km) and the Brecon Beacons (30 cirques in 29 x 7 km) (Fig. 2). With one cirque per 5.2, 6.9 and 6.8 km<sup>2</sup> respectively, each of these sub-regions has a density of cirques rather less than the Maritime Alps (one per 4.0 km<sup>2</sup>). The English Lake District (excluding Black Combe) has 155 cirques in 33 x 31 km (Evans and Cox, 1995, Fig. 1a); one per 6.6 km<sup>2</sup>.

In Wales approximately 35% of cirques are named ‘Cwm...’, although this term is also applied to steep-sided fluvial valleys especially below confluences. The classic cirques of the Snowdon area have been well known since Davis (1909), and mapped and illustrated by Addison (1987). Lewis (1938) considered Llyn Cau on Cadair Idris, and Embleton and Hamann (1988) illustrated Glaslyn on Snowdon, but the cirques in many other parts of Wales have received little attention. Distributions of aspect (azimuth) are summarised by vector mean and vector strength (Evans, 2006), also known as mean direction and mean resultant length. Fig. 3 confirms that most cirques face north or east (Evans, 1999): for all 260 cirques the vector strengths are 53% for axis aspect and 58% for headwall aspect; vector means are  $049^\circ$  for both.

Many, but far from all, cirques were occupied by glaciers in the Devensian Late Glacial, the Loch Lomond Stadial (Gray, 1982; Ballantyne, 2001). Evans (1999) showed that the distribution of cirques occupied by these glaciers or by snowpatches was comparable to that of the whole set, but floors averaged 68 m higher and crests, 90 m higher; larger and better-developed cirques are more likely to have been occupied. For the present study, data have been updated from the work of Lynas (1996), Lowe and Larsen (in Ballantyne, 2001), Carr (2001) and Hughes (2002), mainly extending the Late Glacial occupation of cirques. However, the occupation of 83 cirques remains uncertain, pending considerable further fieldwork. Further references on Welsh moraines and glaciation are given in Evans (1999), Walker and McCarroll (2001) and Lewis and Richards (2005).

[Fig. 3 about here]

## 2.2. *Data variables and definitions*

The main data set used here covers all identified cirques in Wales and supersedes the provisional data set used in a previous study of cirque distribution (Evans, 1999). Cirque form has been re-measured and particular attention has been paid to marginal (debatable) features. One objective is to provide a high-quality data set for comparison with that for the Lake District (Evans and Cox, 1995). The Lake District data have been used by a number of investigators (e.g. Cox, 2004, 2005a) and no measurements have been challenged. The only challenge to the definition of cirques there has come from Wilson (2002) who proposed one extra cirque at Blindtarn Moss west of Grasmere: the headwall had been rejected as relating to a transverse ice flow.

If the development of cirques is to be studied, it is essential to consider all cirques from the most debatable to the most classic and from the smallest to the largest. Hence a complete inventory is needed. Cirques were graded 1 to 5, from classic to marginal or debatable (Fig. 2). Field checks have been extensive, but given the large area over which the 260 cirques are spread in Wales,

checks could not be as thorough as those in the Lake District, and it is likely that detailed investigations will add some further cirques and reject some grade 5 (marginal, debatable) cirques. It is supposed that all definite (grade 3 and better) cirques are included here, and that all measurements are accurate. From the present data, the robustness of results can be tested by repeating analyses with exclusion of marginal cirques, or also of poor ones; or more stringent thresholds for floor or headwall gradient can be set.

Numbers in each grade, 1 to 5, are 23, 51, 68, 59 and 59 respectively. Fig. 4 illustrates the characteristics of differently graded cirques, from less well-known areas. The classic and well-defined cirques (Llyn Lluncaws and Craig Trum y Ddysgl) have sharp contrasts between steep headwalls and flat floors; post-glacial talus accumulation has produced or extended intermediate slopes. Cwm Cwareli has a gentler headwall of alternating sandstone and shale and is less deeply enclosed, but still a definite (average) cirque. Craig Rhiw-erch would also be definite but for the steepness of its floor, which makes it marginal and caused hesitation over its inclusion.

Increasingly, such measurements will be made on-line from scanned maps in GIS, and more automatically by processing DEMs. High-quality DEMs are becoming available and their use is particularly appealing as more broadly-based variables can be defined (Evans, 1987; e.g. gradient-weighted vector mean aspect, and vector strength instead of plan closure). Also profiles and surfaces can be fitted within a cirque. At present, the initial stage of defining and delimiting cirques is subjective and involves air photo interpretation and fieldwork; it takes longer than measurement of the variables used here. The availability of accurate manually measured data sets such as this for Wales should provide useful calibration for future GIS-derived measurements.

Cirques are defined according to the agreed definition reported by Evans and Cox (1974). This is compatible with that of de Martonne (1901). Compared with previous definitions of Snowdonian cirques, the present approach appears to be more stringent in requiring at least part of the headwall to be steep, and more tolerant in accepting sloping floors, up to just over 20°. Thus cirques mapped by both Unwin (1973) and Bennett (1990) at Cwm Bychan (Conwy), Cwm Tŷ-du (Llanberis), Cwm Merch (east of Snowdon), Cwm Planwydd (northeast of Nantlle) and Cwm Ciprwrth (above Cwm Pennant, south of Nantlle) are excluded because they lack cliffs, as is Bennett's Cwm yr Afon Goch (Aber). Their cirque on the northwest slope of Moel Cynghorion, west of Snowdon, is retained in spite of a large landslide from its headwall: before this postglacial event, the cirque floor was lower. Addison (1987, and in Addison et al., 1990, p. 13) was also more demanding than Unwin or Bennett, but omitted most of the Moelwyn-Siabod cirques (which have been modified by the over-riding Welsh ice cap). In the South Wales coalfield numerous landslides are mapped, but their distribution is mainly outside cirques.

[Fig. 4 –photos- about here]

Cirques were classified also by type, most being valley-side (157, including those on escarpments) or valley-head (75, of which two thirds have thresholds). All these are simple, as opposed to nested cirques. Seventeen inner (upper) cirques are contained within 11 outer (lower) cirques: the southernmost are on the north slope of Pumlumon, in the middle of Wales. Ten of the inner cirques are contained in five outer cirques on the Snowdon (Yr Wyddfa) range. This concentration arises because here the summits rise highest above the former glacier equilibrium lines. Furthermore, only on Snowdon itself are three nested levels recognized, for Cwm Llydaw and Cwm Llan; Snowdon has cirques on all sides (Davis, 1909; Addison, 1987; Addison et al., 1990). The criterion for recognizing outer cirques around lower floors is that each should have significant additional headwall that served as an ice source, as opposed to a trough side. When measuring the characteristics of outer cirques, their inner cirques are included except for floor characteristics.

### *2.3. Measurements and their accuracy*

Measurement accuracy depends of course on the care taken by those measuring, and on the quality and resolution (e.g. contour interval) of the maps used. The main differences come from use of different definitions in delimiting cirques, as discussed above. Full definitions of the variables measured and calculated are given in Evans and Cox (1995, Tables 1 and 2); many are in common with those used earlier by Gordon (1977) and by Andrews and Dugdale (1971), facilitating comparisons. Using these tried and tested methods, with a given cirque outline and the accurate photogrammetric contours at 10 m interval on Ordnance Survey 1:10,000 scale maps enlarged for use here, measurements are quite accurate and reproducible. Field experience has increased confidence in the accuracy of these contours.

Once the cirque outline is established it is best next to delimit the floor from the headwall. A spacing of 20 m for 10 m contours gives a slope of 1 in 2, or  $26.6^\circ$ , which usefully defines the boundary between cirque floor (ideally  $<20^\circ$ : Evans and Cox, 1974) and headwall (ideally  $>33^\circ$ , steeper than talus). (A similar spacing is used in delimiting the headwall crest, i.e. the cirque outline, wherever there are gentler slopes above.) Generalizing slightly to give a fairly simple boundary, this permits estimation of maximum floor altitude. It is also needed for locating maximum (head)wall height, along a single slope line, which gives an ancillary measure of vertical dimension.

Outlining the headwall is also useful when it comes to estimating the cirque focal point, the mid-point of the threshold or sill. This is straightforward for cirques which are internally symmetrical, but many cirques deviate from this and many thresholds are sloping, so the mid-point deviates from the lowest point. In these cases there is an inevitable subjectivity, as it is necessary to compromise between the mid-points given by several contours on the headwall, and that half-way between the intersections of the headwall-floor boundary with the cirque outline. Differences in the focal point shift the median axis, changing its aspect and length.

The median axis (Unwin, 1973) has been visually estimated as leaving half the cirque map area to the left, and half to the right. Use of tracings to superimpose the two sides suggests that initial visual estimates give axial aspects within  $5^\circ$  of the final value, which in turn is within  $2^\circ$  of the true value. Thus such manual estimates are fully comparable with the exact measurements now possible in GIS (Federici and Spagnolo, 2004). Estimates of headwall aspect are less precise, but they are unaffected by uncertainty over the focal point and less affected by variations in cirque outline definition, such as those due to possible headwall extensions with marginal gradients. Here the future use of DEMs will permit calculation of a headwall resultant vector based on point values of slope aspect, weighted by gradient in excess of the  $26.6^\circ$  threshold. The present manual estimates are considered accurate within  $10^\circ$ . It is reassuring also that the difference between headwall and axis aspect is symmetrically distributed, with a standard deviation of only  $20^\circ$ ; its range is  $\pm 65^\circ$ .

Six altitude variables are defined. Estimates are accurate within 5 m, but considerable variance may come from differences in cirque outline, floor boundary and median axis. Thus the most reproducible variables are probably modal floor altitude and maximum crest altitude. The maximum altitude above, draining into a cirque, is often given exactly by a spot height, but estimates could be out by tens of metres where the location is uncertain on a mountain shoulder with divergent flow.

Again given the focal point and outline, length and width measurements are accurate within 10 m. Larger differences come from shifting the focal point or changing the cirque boundary. Visual estimates of the perpendicular nature of length and width were within  $2^\circ$ . Plan closure (Evans, 1969; Gordon, 1977) is a variable that requires some experience, to avoid problems with quadrants and complementary angles, but after careful measurement reproducibility within a few degrees is achieved.

With these methods of measurement and estimation from maps, and given a map with the cirque outline, all directly measured variables can be established within about an hour. This varies with cirque size, between about 30 and 90 minutes. Once a data set is complete and entered into a

statistical program, hours should be spent on consistency and outlier checks, extending to days if numerous errors are discovered and corrected. Before the final measurement by the author, most of these cirques were measured at different times, from different maps, and/or by different people, under the author's supervision. This improved the establishment of the best cirque outline, and aided avoidance of substantial errors.

[Fig. 5 about here]

### 3. Results

#### 3.1. Size

Welsh cirques are comparable in size to those described from Scotland and the English Lake District. They do not fall into discrete size classes, but give fairly smooth continua on all measures of size. Positive skewness ranges from 1.22 for wall height to 2.45 for length. After logarithmic transformations these are reduced to  $-0.06$  and  $0.32$ . For variables measuring components of vertical dimension, generated as differences between altitude variables, skewness of  $0.76$  to  $2.49$  is reduced to  $-0.22$  to  $-0.76$ . Fig. 5a shows how well-behaved the main size variables are on a logarithmic scale. These quantile plots (Cleveland, 1994, pp. 143-9) permit the observed distribution to be compared with a model probability distribution, in this case the log-Gaussian (also known as log-normal). Compared with histograms, they emphasize the overall shape and show every observation, unaffected by binning (classing), and they 'stack' onto one plot to facilitate comparisons (Cox, 2005b). Their linearity confirms that each is well approximated by the log-Gaussian model, and thus all further analyses deal with logarithms of these size variables. Note that each variable is ranked separately, so vertical comparisons on Fig. 5a generally concern different cases.

As is usual, extreme values are more erratic than those in the middles of the distributions. One notable effect is that ranked length is less than ranked width except for the four longest cirques. These are in fact the outer cirques of Cwms Llydaw, Llan and Dwythwch, all on Snowdon, plus Cwmffynnon on the Glyderau. (Except for Cwm Dwythwch these are also the three widest cirques, although the ranking process for the plots is independent for each variable.) Omission of all 11 'outer' cirques reduces skewness, but when relations between variables are considered (below) these cirques are 'on trend' and strengthen rather than weaken correlations, so it is best to retain them.

As in other regions, for example the Rocky Mountains of the USA (Graf, 1976) or Canada (Trenhaile, 1976), Welsh cirques have width close to length so that plan form is compact. Median

(and mean) values for Wales are 610 (667) m for length, 700 (772) m for width and 215 (236) m for amplitude. Spreads represented by the 5 and 95 percentiles (used here in preference to the range of extreme values) are 310-1235 m for length, 375-1502 m for width and 112-431 m for amplitude. These dimensions are comparable to those in England and in Scotland (Gordon, 1977), but rather smaller than those in northern Scandinavia (Hassinen, 1998) and considerably smaller than those in the Canadian Rockies and Columbia Mountains (Trenhaile, 1976) and the Antarctic Dry Valleys (Aniya and Welch, 1981). Cirques in the Maritime Alps (Federici and Spagnolo, 2004), the central Spanish Pyrenees (García-Ruiz et al., 2000) and the Cayoosh Range of British Columbia (Evans, 1994, Table 26.3) are comparable to Welsh cirques in area, but have rather greater vertical dimensions.

Confidence intervals show that cirque width is significantly greater than length, more so in Wales than in the Lake District. Width exceeds length also in the central Pyrenees (García-Ruiz et al., 2000), in central Sweden (Vilborg, 1984) and in northern Scandinavia (Hassinen, 1998) but in most other studies (northwest Scotland, the Maritime Alps, the whole Italian Piemonte, the Cayoosh Range, Baffin Island, the Canadian Rockies and Columbia Mountains and the Antarctic Dry Valleys) length is commonly greater. Width/length ratio averages 1.22 in Wales, and ranges from 0.49 to 3.47, so only one cirque is twice as long as broad. Cirques with width over twice length tend to be poor or marginal. Six of the 16 are in the Rhondda – Hirwaun group, and a further six are also in South Wales; all but one of these are on sandstone.

[Fig. 6 about here]

Fig. 6 shows, in five equal classes, the spatial distribution of cirque (overall) size, a combination of the three orthogonal dimensions. Size is defined as the cube root of (length x width x amplitude), so that it too has units of metres: it is required for analyses later in the paper. Larger cirques are concentrated in northern Snowdonia, and smaller ones are more common in South Wales. Cadair Idris and most northern regions have means above 460 m while the Brecon Beacons average 448 m and Corris and the rest of South Wales have means below 420 m, except that Pumlumon (mean 486 m) can compete with the northern regions and Migneint and Upper Dyfi bring down their regional group means.

These regional variations in the size and shape of Welsh cirques are further demonstrated by the variables in Figs. 7 and 8. The point-and-box plots show both the individual data and the quartile summaries (25, 50 and 75 percentiles). They are preferred over conventional box plots,



which can mislead by masking the number of cases on which they are based, and which have difficulties handling log scales because inter-quartile range is dependent on the metric.

As expected, the greatest vertical dimensions are in the Snowdon and Glyder groups, followed by Carneddau, Nantlle – Hebog, Cadair Idris and the Aran Range (Fig. 1), with mean amplitudes greater than 250 m. All these are within the main Ordovician volcanic belt, although some cirques are on intercalated siltstones. The least deep vertically are in South Wales (means less than 195 m) and in the Migneint area (between the Arenig and Moelwyn mountains), followed by Corris and Pumlumon (Fig. 7a). These are on sedimentary or lightly metamorphosed rocks. The low outlier for Arenig – Migneint is Llynau Barlwyd, where the median axis intersects a very low central col, an unusual situation.

Length too is greater in North Wales than in the South (Fig. 7b) but Cadair Idris, Nantlle – Hebog and the Carneddau have surprisingly low medians and the cirques of Mid Wales (Pumlumon) are surprisingly long. There is, however, much less North – South contrast in width (not illustrated), which is greatest in the Western Brecon Beacons and Moelwyn – Siabod and least in the Black Mountains, Abergavenny and Corris.

These results are affected by inclusion of marginal cirques, which form 30% of the total from Corris southward, but only 18% farther north. Overall, in spite of some impressive cirques on sandstone in South Wales, cirques on sedimentary, slaty and greywacke rocks are smaller than those on volcanic and igneous rocks. For analyses of variance over the 17 regional groups,  $R^2$  values are 0.294, 0.093 and 0.014 for the logarithms of amplitude, length and width respectively. Length is significant at  $P=0.0007$ , whereas width is quite insignificant.

Fig. 5b shows that altitude variables are well distributed without transformation, across broad ranges, except that ‘high’ altitudes cannot exceed the limiting value of 1065 m, the highest point in Wales.

[Fig. 7 about here]

### 3.2. Shape and gradient

Gradient and closure variables are symmetrically distributed (Fig. 5c and d) except for minimum gradient where proximity to the lower bounding value of zero inevitably produces positive skew. Reversed slopes in lakes and bogs are excluded because of inadequate data. As in the Lake District (Evans and Cox, 1995, Fig. 5), quantile plots for plan closure (Fig. 5d) and axial gradient are very linear, confirming well-behaved Gaussian distributions with no need for

transformation. Profile closure and maximum and minimum gradient have tails shorter than the Gaussian, giving mildly S-shaped quantile plots.

Measuring how deeply cirques cut into mountains, closure in plan is greatest (mean  $149^\circ$ ) in the Pumlumon group, where most cirques are at the heads of valleys incised into the plateau. Otherwise it is greater in groups containing igneous rocks, including the Berwyns (Fig. 7c). The high outlier for Moelwyn is Cwmorthin ( $295^\circ$ ), which has a complicated headwall. Easily the lowest plan closures are in the Black Mountains ( $61^\circ$ ), where the four cirques are shallow recesses in long, smooth valley sides. Other sandstone groups, including the Rhinogs, have relatively low closures. Over the 17 regional groups,  $R^2$  is 0.073,  $P = 0.0043$ .

Maximum gradients (not illustrated) show a clear contrast between the five mountain groups of northern Snowdonia, plus Cadair Idris and the Arans, and all the rest. Low values are found south of Cadair Idris, especially in the Western Brecon Beacons. Minimum gradients are lowest throughout Snowdonia. High minima for Corris (mean  $9.4^\circ$ ) and the six Upper Dyfi cirques reflect the inclusion of marginal cirques on the weak Silurian and Upper Ordovician siltstones.  $R^2$  is 0.191 for maximum and 0.158 for minimum gradient; both are highly significant.

Axial gradient (Fig. 7d;  $R^2$  is 0.180) is a more representative measure of overall cirque steepness. All groups have values both above and below  $20^\circ$ , but Corris and the four north-western Snowdonia groups have means above  $21^\circ$ . Moelwyn – Siabod joins the three groups immediately to its south in having low gradients, together with all of South Wales where the Brecon Beacons have means just below  $18^\circ$ . This mixed picture arises because cirques are steep both where relief is highest (around Snowdon) and where cirque development is more marginal (as around Corris).

As a glacier filling a cirque is likely to start near the top of the median axis, axial gradients are usefully compared with the gradients of cirque glaciers. Glaciers not reaching that high up the headwall, and those overflowing the cirque threshold, would have gentler gradients. Only 10% of Welsh axial gradients are less than  $13^\circ$ , which is sufficient to support rotational flow (section 1.2). Only three (1% of cirques) are gentler than  $8^\circ$ ; 10% are steeper than  $30^\circ$ , and the maximum axial gradient is  $38^\circ$ . Thus the great majority of cirques can be related to cirque glaciers or glacier sources capable of rotational flow.

As in the Lake District, correlations within this group of shape and gradient variables are weak, except that axial gradient correlates +0.63 with minimum gradient. Maximum and minimum gradients correlate only  $-0.26$ . Plan closure correlates  $-0.35$  with axial, but only  $\pm 0.19$  with the other two gradients (still significant, at the 0.005 level), and  $-0.20$  with width/length. Principal component analysis of six shape and gradient variables confirms the weakness of their interrelations; the components have 40, 26, 19, 12, 3 and 0 percent of the total variance. This

contrasts with the six altitude variables, for which the first component has 86%, and the seven logarithmic size variables, 71%.

### 3.3. Geological effects

All erosional landforms are affected by the material into which they cut, but it is often difficult to pin down precise morphometric contrasts between lithologies (Evans, 1994). In Snowdonia, Unwin (1973, p. 87) noted structural control of detail but “considerable disregard for geological structure” with cirques cross-cutting different lithologies in the Nantlle – Hebog group. In the Snowdon group and especially in Y Glyderau, some cirques are elongated along the strike. Floors are often on weaker strata, giving exaggerated forms deep in profile in several cirques of Y Carneddau. Bennett (1990) developed the concept of ‘strike cirques’ further.

The importance of joints, faults and other planes of parting in cirque development was emphasized for example by Haynes (1968) and Addison (1981). In Snowdonia, the fracture network disregards lithological boundaries, but lithology controls fracture spacing and rock mass strength at smaller scales (Addison, 1981). However, data are not available on a broad basis and for the present survey the units mapped by the British Geological Survey and others are used. Thirteen units are distinguished in Fig. 8; 22% of Welsh cirques are on Devonian and Carboniferous sandstones (with intercalated shales), 14% are on weak Ordovician and Silurian siltstones and greywackes, 19% are on tuffaceous Ordovician sediments, 29% are on Ordovician volcanic rocks, 8% are on Cambrian rocks and 6% are on intrusive granitic rocks and dolerites.

For each variable, a one-way analysis of variance gives the variability accounted for by the 13 classes of geology. Table 1 ranks the 14 variables, and shows that the most closely related to geology are vertical dimensions and altitudes, followed by gradients (especially maximum) and plan closure. The relations are weaker than those over regional groups, but their ranking is comparable. The variation of length with geology is small (Fig. 8a), and that of width is insignificant. Coal Measures cirques (which have mainly Pennant sandstone headwalls) are shorter than others, and have lower amplitudes, but are just as wide. Lengths are greatest on tuffs.

[Table 1 about here]

The strongest relation to geology comes from relief (within 2 km radius), which is another controlling factor rather than a cirque characteristic. This suggests considering the joint effects of the two on the other variables. Relief is in fact the greater control for maximum gradient, axial gradient, size and length, with no additional effect from geology. It is also the greater for maximum

crest altitude, wall height, height range and amplitude, for which geology has small but significant further effects. On the other hand, relief is quite unrelated to plan closure, width and minimum gradient, and has no significant addition for lowest altitude. Thus it seems that for vertical dimensions, related gradients, length and thus size, relief is the direct control, but is itself affected by geology. Geology affects plan closure and minimum gradient (floor development) directly. For a selection of well-developed Austrian and British cirques, Embleton and Hamann (1988) also found that relief was more important than geology in controlling cirque form.

Rock basin lakes are a characteristic more closely related to geology (Haynes, 1968; Evans, 1994). In Wales, 11 of the 21 major rock basin lakes are on 'ignimbrite, lava and rhyolite', forming 29% of the cirques on that geological class. There are none on Lower Ordovician siltstones or on Silurian or younger rocks, although there are three major moraine-dammed lakes on Coal Measures sandstones and three more on Old Red Sandstone (Devonian).

[Fig. 8 about here]

Fig. 8b shows greater cirque amplitude (height range along the median axis) for volcanic and igneous rocks and tuffaceous sediments than for other rock types; amplitude is lowest on Devonian and Carboniferous rocks. As amplitude is more variable between rock types than are length and width (Table 1), variation in the compound variable 'size' is similar to that for amplitude. Relief (Fig. 8c) shows a similar pattern (and relief at 1 km radius correlates +0.89 with relief at 2 km), with clearly lower values on the younger, sedimentary rocks of the South, including the Silurian, and highest values on the three volcanic rock types and on microgranite; tuffaceous siltstones rank lower than for amplitude.

Maximum gradient (Fig. 8d) shows a clear dichotomy around  $64^\circ$ , between steep headwalls on volcanic and igneous rocks and gentle ones on sedimentary rocks – including Cambrian grits, and tuffaceous siltstones but not tuffaceous sandstones. Plan closure shows a different pattern (Fig. 8e), being greater on Silurian and Lower Ordovician siltstones, mudstones and greywackes (with numerous valley-head cirques), and lowest on the intrusive rocks (microgranite and dolerite) which are difficult to incise. It is fairly low on sandstones: around  $100^\circ$ , whether Cambrian, Devonian or Carboniferous.

### *3.4. Effects of other controls*

As expected, the altitude and size groups of variables each have strong intercorrelations. The three 'high' altitude variables all correlate +0.91 or more: the three 'low' ones, +0.93 or more,

and all six, +0.68 or more. Width and length, as logarithms, correlate +0.72 . Length correlates +0.62 to +0.71 with the three measures of vertical dimension, whereas width correlates only +0.36 to +0.45.

Thus it is necessary to be selective in relating characteristics to combinations of others. Length, width and amplitude all increase with (any of the three) maximum altitudes, and they increase as lowest altitude decreases. Regressions including dummy variables for categories show that length is a function first of maximum altitude above, followed by lowest altitude, cirque type and occupation, with  $R^2 = 0.54$ , to which relief adds a little but geology adds nothing. Width gives  $R^2 = 0.36$  for the same four, with lowest altitude being more important than maximum: relief increases this to 0.41. Amplitude relates more to maximum altitude than to lowest, and as it is calculated as the difference between a high and a low variable, including both gives a spuriously strong prediction.  $R^2 = 0.41$  for amplitude as a function of maximum altitude above, cirque type and geology; relief increases this to 0.47 but occupation gives no further improvement.

[Table 2 about here]

#### 4. Allometry

##### 4.1. *Variations of shape with size*

Table 2 shows that maximum headwall gradient increases with all five size variables, and minimum floor gradient decreases. But maximum increases mainly with vertical dimensions, to which minimum is barely related. Overall axial gradient, as expected, rises with vertical dimensions and falls with increasing length and width. The width/length ratio has modest correlations with all size variables; these are negative for vertical dimensions, which as noted above relate more to length than to width. Plan closure also relates first to length, then to width. Larger cirques are likely to have more cols (over 30 m deep) in their crests; average sizes are 440 m for no cols, 566 m for one, 743 m for two, 754 m for three, and 1501 m for the one cirque with four cols.

[Fig. 9 about here]

A more graphic way of relating shape to size is to divide size into five equal classes and draw representative profiles for each (Fig. 9). The choice of five classes gives a reasonable number, 52, in each class for reliable estimation of mean values. For each size class, modal and maximum floor altitude, median and maximum crest altitude, and maximum altitude above are plotted relative to lowest altitude, and against horizontal coordinates based on length, minimum gradient (lowest

segment), maximum gradient (either side of the dot), and an assumed  $10^\circ$  above the crest. This is comparable to the developmental diagrams in Gordon (1977) and Bennett (1990), but based more directly on the data. Assuming a space-time transformation, headwall recession is greater than vertical enlargement, but the latter is considerable.

Development in length and width can be portrayed by a similar series (Fig. 10), using mean plan closures to dictate the curvatures. The full width is used, but as plan closure is measured along the mid-height contour which does not extend to either end of the median axis, only half the length is represented here: these are not cirque outlines. Shape changes are more subtle than size changes, but there is a progressive broadening out.

[Fig. 10 about here]

#### *4.2. Allometric development*

The concept of allometry, widely applied in developmental biology, simply implies that shape changes as organisms grow. Thus it is opposed to isometry, the maintenance of constant shape. These alternatives are commonly assessed on logarithmic scales, so that a given interval on an axis represents multiplication by a given factor. Applications to surveys of landforms at a single time are based on the assumption that the landforms have grown over time. This is not unreasonable for bedrock landforms, especially those such as cirques where positive feedback is involved. Application of allometry to glacial cirques was initiated by Olyphant (1981, p. 681), who took 'volume' (length x width x depth) as a reference measure of size to which other size measures could be related. He defined length as from cirque lip to headwall midpoint, width as the average of four measurements equally spaced along the length axis, starting at the cirque lip, and depth as the vertical difference between the cirque lip and the average cirque rim altitude. The width and depth measures are more averaged than those used here, but this is likely to affect the base constants rather than the gradients of logarithmic relations.

Here, following Evans and Cox (1995) and Evans and McClean (1995), I use median-axial length, maximum width at right angles to that axis, and vertical amplitude as the difference between crest altitude at the median axis and lowest altitude. 'Volume' is the product of the three linear dimensions, and 'size' is its cube root, useful in giving a linear measure comparable in magnitude to the three original measures. As  $\text{volume} = \text{size}^3$ , the power exponents (gradients of logarithmic regressions) are three times as great for size as for volume. That is, while length, width and amplitude exponents sum to 1.0 for predictions from volume, they sum to 3.0 for predictions from size.

An advantage of the present data set is that there are two further definitions of vertical dimension, whose exponents can be compared with those for amplitude. Wall height is measured independently as the greatest drop along any headwall slope line, from crest to floor at right angles to contours. Overall height range is the extreme, relating maximum crest altitude to lowest altitude: it cannot be less than amplitude or wall height. It is also possible to compare exponents for different geologies or grades. Confidence intervals are considered the best way of comparing the feasible variations in estimates of exponents, and the significance of differences between them.

[Table 3 about here]

Table 3 confirms that exponents for the three components of size do indeed sum to 3.0. It shows that length increases (with size) faster than does width, and width probably increases faster than amplitude. Width and wall height show exponents not significantly different from 1.0 throughout; thus they are isometric, increasing in proportion to overall size. Length has significantly greater exponents; it increases faster than size. The confidence interval on the length exponent does not overlap with any other, except wall height. The results for overall height range are very similar to those for amplitude, except for valley-head cirques. The ‘better’ column, where poor and debatable cirques are excluded, gives very similar exponents except for wall height. Although ‘outer’ cirques form a distinctly larger population, their exclusion makes no difference to the exponents obtained. Valley-side cirques show almost the same exponents as the total set, but the smaller set of valley-head cirques does give a steeper variation in amplitude with size, and a reduced variation in length. Fig. 11 shows that the scatter of data points around these regressions is well-behaved, with near constant variance on a logarithmic scale.

[Fig. 11 about here]

These results provide a powerful confirmation of cirque allometry, robust over different cirque types. Length increases faster than size, and vertical dimension increases more slowly whether axial or overall height range is used (Fig. 11); this is consistent with Gordon’s (1977, Fig. 7) model. Large cirques are relatively longer and less deep than if growth were isometric. The isometry of wall height is interesting, but it does show considerably more scatter than the other vertical variables in both dimensionless ( $R^2$ ) and dimensional (RMS error) terms.

[Table 4 about here]

From Table 4 it seems that allometry of cirques in Wales is less pronounced than in the other areas. This may reflect the varied relief and geology of Wales. The greater increase of length than width is shared with the English Lake District and the Cayoosh Range. It is contrary to what would occur if allometry were due to lateral coalescence, which would increase width rather than depth.

Relations within different regions and broad rock types will now be considered. As confidence intervals are too broad for the 13 classes, three broader divisions of geology are used: igneous and volcanic; other, that is Cambrian and Ordovician sediments including tuffaceous; and Silurian and younger rocks. In relation to Fig. 8, these three are defined as the first five geological classes, the next four, and the final four which equate to South Wales and parts of Mid Wales.

In Table 5, the strongly overlapping confidence intervals for each variable show that none has exponents differing significantly between these three rock types. The high  $R^2$  and low RMS errors confirm the strengths of these relations. The poorer result for amplitude on 'other' rocks is due to the outlying but real value for Llynau Barlwyd (on siltstones), noted above: thus overall height range gives a better result for 'other'. Wall height shows greater scatter than the other variables, and unlike height range and amplitude its exponents for igneous and other do not differ significantly from 1.0.

[Table 5 about here]

Throughout Table 5, length exponents consistently exceed those for width, except on the Silurian and younger rocks, and all vertical dimensions produce lower exponents. Thus the results of Table 3 apply across the three main rock divisions. A similar analysis (not shown) for three broad regional divisions, North (Carneddau to Moelwyn-Siabod), Mid (Arenig-Migneint to Pumlumon) and South (Brecon Beacons to Rhondda) further reinforced the results. Here all three vertical dimensions gave exponent values significantly less than 1.0 for North and South, but closer to 1.0 for Mid Wales, which mixes high and low relief. Length and width gave highly overlapping confidence intervals, with the length exponent greater in North but the width exponent slightly greater in South.

## 5. Summary and Conclusions

In Wales, width exceeds length in 66% of the 260 cirques. This ratio applies across different cirque types, but rises to 91% for the 56 cirques on Devonian and Carboniferous sandstones, often on escarpments or straight valley-sides. Width averages 772 m, length 667 m,



(axial) amplitude 236 m, overall height range 269 m and wall height 197 m. Cirque dimensions are thus comparable to those in other mid-latitude studies, and especially to those in the English Lake District. Axial gradients are compatible with rotational flow of former glaciers in the great majority of these cirques.

Although differences between regions in length and width and their variations could arise from different practices of cirque definition, this is not the case for the contrast between Wales and the Lake District, with mean length/width ratios of 0.90 and 0.94, and the Cayoosh Range with 1.14, as I am responsible for all these definitions. The differences probably reflect regional topographic and tectonic settings, especially the degree of dissection of the landscape and the available relief. Although length and width behave somewhat differently, the slower increase of vertical dimension with cirque size is observed in all these regions.

If size (a combination of length, width and amplitude) is divided into five quintiles, each of its components (and the other vertical dimensions) increases monotonically, but there are changes in shape. Length increases most, vertical dimensions least. Plan and profile closure increase, and maximum gradient increases faster than minimum decreases.

These data provide the strongest test yet of allometry in glacial cirques. As elsewhere, vertical dimensions – amplitude, height range and wall height – increase with overall size, but significantly less than do horizontal dimensions. In Wales, only length has an exponent significantly greater than 1.0; that is, it increases faster than size in general. With minor variations, these results hold over different cirque types, regions and rock types. This type of allometry is logical given the role of glacial erosion in cirque development; the glacier rises higher against the backwall and moves away from it faster than from the sidewalls, thus increasing length faster than width. It is also reasonable that headwall retreat, often by collapse following glacial erosion at the base (Evans, 1997), is faster than downward erosion. Such importance of cirque headwall retreat has been emphasized recently by Mitchell and Montgomery (2006) and especially by Oskin and Burbank (2005), raising the question whether glacial erosion in mid-latitude mountains works largely as a ‘buzzsaw’ at the snowline or mainly by the calibration of glacial troughs to ice discharge. The divide retreat proposed by Evans (1972), Oskin and Burbank (2005) and others, shows that lateral coalescence of cirques cannot explain the greater increase of horizontal dimensions, especially since length increases with size faster than does width.

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## Figures

Fig. 1. Wales: altitudes of the highest summits in each of the 17 cirque groups used here (e.g. Fig. 7), plus important secondary summits. Names of the groups are used (or parts of compound names such as Moelwyn – Siabod and Nantlle – Hebog), rather than summit names. Here, and on Figs. 2 and 3, the grid lines are kilometres on the Ordnance Survey National Grid, a variant of the Universal Transverse Mercator.

Fig. 2. The distribution of cirques in Wales, by grade.

Grades are defined as:

1. Classic, with all textbook attributes, – a steep headwall curving around a deeply excavated floor;
2. Well-defined, with the headwall curving around the floor and both clearly developed;
3. Definite, with no debate about cirque status but one weak characteristic;
4. Poor, open to some doubt but with better characteristics compensating for weaker ones; and
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Fig. 3. The distribution and headwall aspects of cirques in three parts of Wales by size for (a): Northern Snowdonia, (b): the Dolgellau region (Meirionnydd district), and (c): South Wales. The scales vary, and are shown by the grid coordinates in km. Arrows start from cirque mid-points and their length is proportional to cirque size.

Fig. 4. Photographs of cirques of various grades of development.

a: Llyn Lluncaws, in a classic cirque in the Berwyns.

b: Craig Trum y Ddysgl, the west side of a well-defined cirque above the Nantlle, viewed from Mynydd Drws-y-coed.

c: Cwm Cwareli, a definite cirque near the east end of the Brecon Beacons.

d: Craig Rhiw-erch, a cirque on Maesglase at the northeast end of the Corris group; it is marginal because of its steep (although drift-filled) floor.

Fig. 5. Quantile plots of measured variables, each individually in rank order, for (a): size, (b): altitude, (c): gradient and profile closure, and (d): plan closure.

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**Table 1**

Analysis of variance results, in rank order, for relation of cirque characteristics to the 13 classes of geology in Fig.8

| <i>Variable</i>     | <i>adjusted <math>R^2</math></i> | <i>variance ratio, <math>F</math></i> | <i>P-value</i> |
|---------------------|----------------------------------|---------------------------------------|----------------|
| log(Relief in 2 km) | .506                             | 23.1                                  | .0000          |
| log(Relief in 1 km) | .438                             | 17.8                                  | .0000          |
| Max. crest altitude | .340                             | 12.1                                  | .0000          |
| log(Wall height)    | .250                             | 8.2                                   | .0000          |
| log(Height range)   | .234                             | 7.6                                   | .0000          |
| Lowest altitude     | .208                             | 6.7                                   | .0000          |
| log(Amplitude)      | .199                             | 6.4                                   | .0000          |
| Max. gradient       | .165                             | 5.3                                   | .0000          |
| Plan closure        | .083                             | 3.0                                   | .0007          |
| Min. gradient       | .054                             | 2.2                                   | .0113          |
| log(Size)           | .046                             | 2.0                                   | .0220          |
| Axial gradient      | .045                             | 2.0                                   | .0231          |
| log(Length)         | .038                             | 1.9                                   | .0398          |
| log(Width)          | -.020                            | 0.6                                   | .8688          |

Note that adjustment for the 12 fitted constants has reduced  $R^2$  values considerably.

**Table 2**

100x correlations of five size variables with five shape or gradient variables

|                        | <i>max. gradient</i> | <i>min. gradient</i> | <i>axial gradient</i> | <i>plan closure</i> | <i>width/length</i> |
|------------------------|----------------------|----------------------|-----------------------|---------------------|---------------------|
| <i>log length</i>      | 38                   | -51                  | -49                   | 51                  | -38                 |
| <i>log width</i>       | 30                   | -55                  | -48                   | 39                  | 34                  |
| <i>log amplitude</i>   | 53                   | -00                  | 36                    | 22                  | -35                 |
| <i>log heightrange</i> | 58                   | -12                  | 20                    | 33                  | -35                 |
| <i>log wall height</i> | 57                   | -20                  | 23                    | 29                  | -33                 |

**Table 3**

Exponents for logarithmic (power) regressions of size variables on overall cirque size in Wales

| <i>Variable</i> | <i>expon.</i> | <i>95% conf.</i> | $R^2$ , % | <i>st.dev.</i> |               |                 |                    |                    |
|-----------------|---------------|------------------|-----------|----------------|---------------|-----------------|--------------------|--------------------|
|                 |               |                  |           |                | <i>better</i> | <i>no outer</i> | <i>valley-side</i> | <i>valley-head</i> |
| Length          | 1.12          | 1.07-1.18        | 86        | 0.16           | 1.10          | 1.12            | 1.13               | 1.01               |
| Width           | 0.98          | 0.89-1.06        | 68        | 0.17           | 0.98          | 0.99            | 0.97               | 0.94               |
| Amplitude       | 0.90          | 0.81-0.99        | 61        | 0.16           | 0.91          | 0.89            | 0.90               | 1.05               |
| Height range    | 0.91          | 0.83-0.99        | 67        | 0.15           | 0.89          | 0.90            | 0.90               | 0.93               |
| Wall height     | 0.97          | 0.86-1.09        | 52        | 0.19           | 0.85          | 0.97            | 1.02               | 0.99               |

95% confidence intervals and  $R^2$  measures of fit for all 260 cirques in Wales are given on the left. These are followed by the standard deviation of each variable, and exponents for 142 better cirques (graded definite, well-defined or classic), for the 249 cirques excluding 'outer' cirques, for 157 valley-side and for 75 valley-head cirques

**Table 4**

Comparative exponents for logarithmic (power) regressions of size variables on overall cirque size

| <i>Variable</i>          | <i>Wales</i> | <i>Lake D.</i> | <i>Cayoosh</i> | <i>Mar. Alps</i> | <i>Blanca</i> |
|--------------------------|--------------|----------------|----------------|------------------|---------------|
| <i>number of cirques</i> | 260          | 158            | 198            | 432              | 15            |
| Length                   | 1.12         | 1.17           | 1.10           | 1.08             | 1.14          |
| Width                    | 0.98         | 1.10           | 1.05           | 1.08             | 1.20          |
| Amplitude                | 0.90         | 0.74           | 0.85           |                  |               |
| Height range             | 0.91         | 0.75           | 0.83           | 0.84             | 0.66          |

Results for the whole of Wales are compared with those for the English Lake District (Evans and Cox, 1995, p. 195), the Cayoosh Range of British Columbia (Evans and McClean, 1995, p. 136), the Maritime Alps of Italy and France (Federici and Spagnolo, 2004, p. 240), and the Blanca Massif in the southern Colorado Rockies. The two former are in 'old massifs': the others, in active orogenic belts.

**Table 5**

Exponents, with 95% confidence intervals, for logarithmic (power) regressions of size variables on overall size, for three mapped geologies

| <i>Variable</i> | <i>geology</i> | <i>expon.</i> | <i>95% conf.</i> | <i>R<sup>2</sup>, %</i> | <i>RMS error</i> |
|-----------------|----------------|---------------|------------------|-------------------------|------------------|
| Length          | igneous        | 1.11          | 1.04-1.18        | 92                      | .054             |
|                 | other          | 1.18          | 1.06-1.30        | 81                      | .077             |
|                 | younger        | 1.10          | 0.98-1.21        | 83                      | .067             |
| Width           | igneous        | 1.04          | 0.93-1.16        | 78                      | .093             |
|                 | other          | 0.96          | 0.82-1.10        | 66                      | .093             |
|                 | younger        | 1.13          | 0.97-1.29        | 73                      | .093             |
| Amplitude       | igneous        | 0.85          | 0.71-0.96        | 65                      | .104             |
|                 | other          | 0.87          | 0.68-1.05        | 48                      | .122             |
|                 | younger        | 0.77          | 0.66-0.87        | 70                      | .067             |
| Height range    | igneous        | 0.87          | 0.74-0.99        | 67                      | .101             |
|                 | other          | 0.82          | 0.69-0.95        | 64                      | .083             |
|                 | younger        | 0.83          | 0.71-0.95        | 71                      | .071             |
| Wall height     | igneous        | 0.87          | 0.69-1.05        | 51                      | .141             |
|                 | other          | 0.91          | 0.72-1.10        | 50                      | .124             |
|                 | younger        | 0.82          | 0.64-1.00        | 51                      | .107             |

‘Igneous’ are intrusive and volcanic (91 cirques), ‘other’ are Cambrian and Ordovician sediments including tuffaceous (93 cirques), and ‘younger’ are Silurian, Devonian and Carboniferous rocks (76 cirques).

# **FIGURES for “Allometric development of glacial cirque form: geological, relief and regional effects on the cirques of Wales”**

**Ian S. Evans**

(submitted 28 Aug 2005, revised 19 January & corrected 22 February 2005)

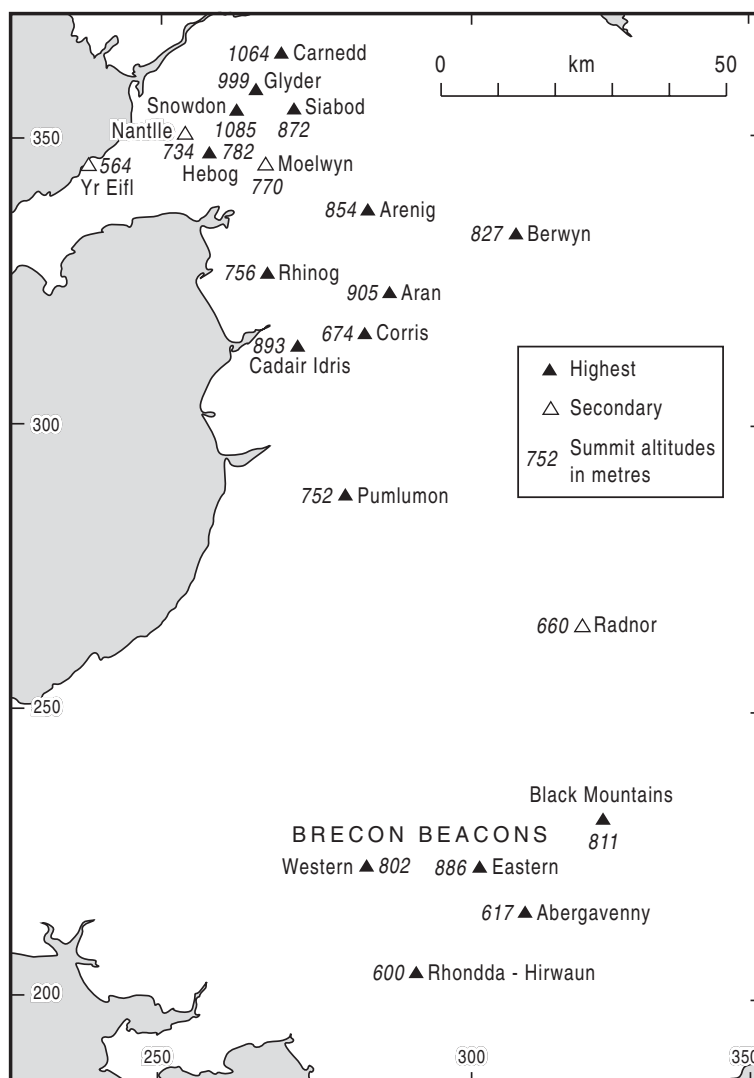


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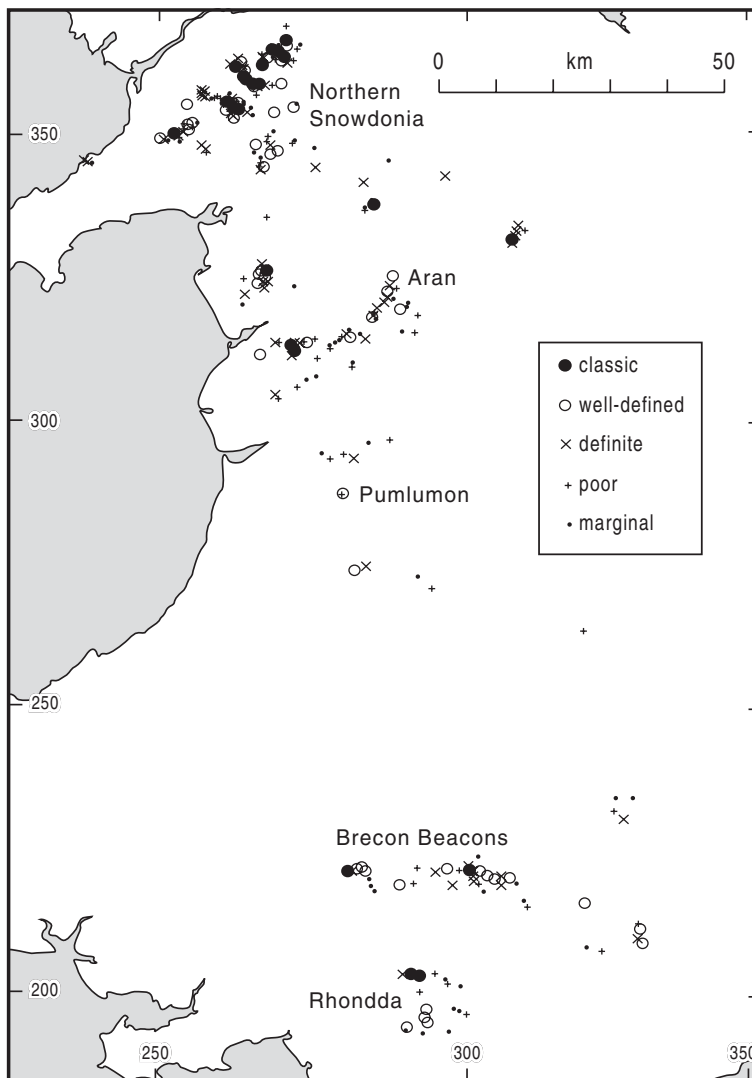


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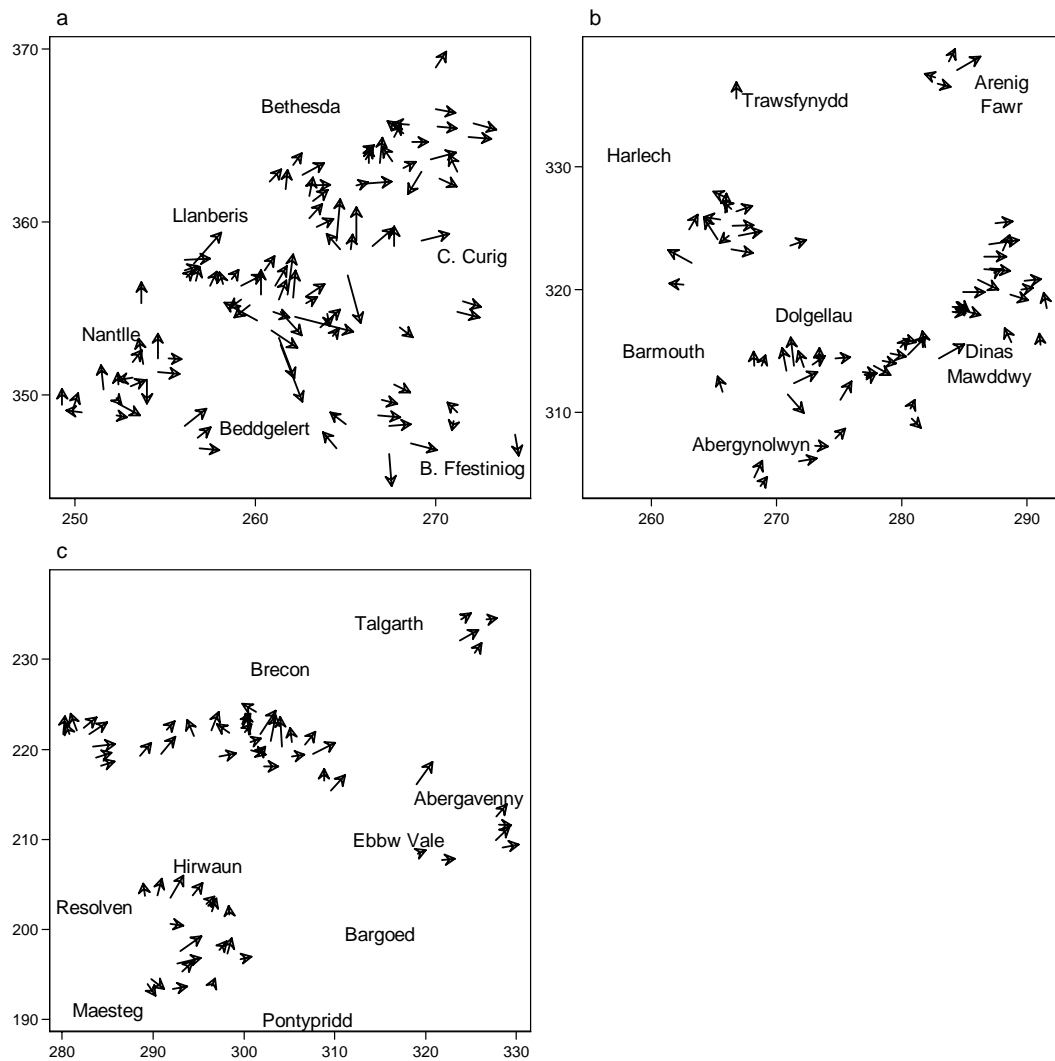


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a:



b:



c:



d:

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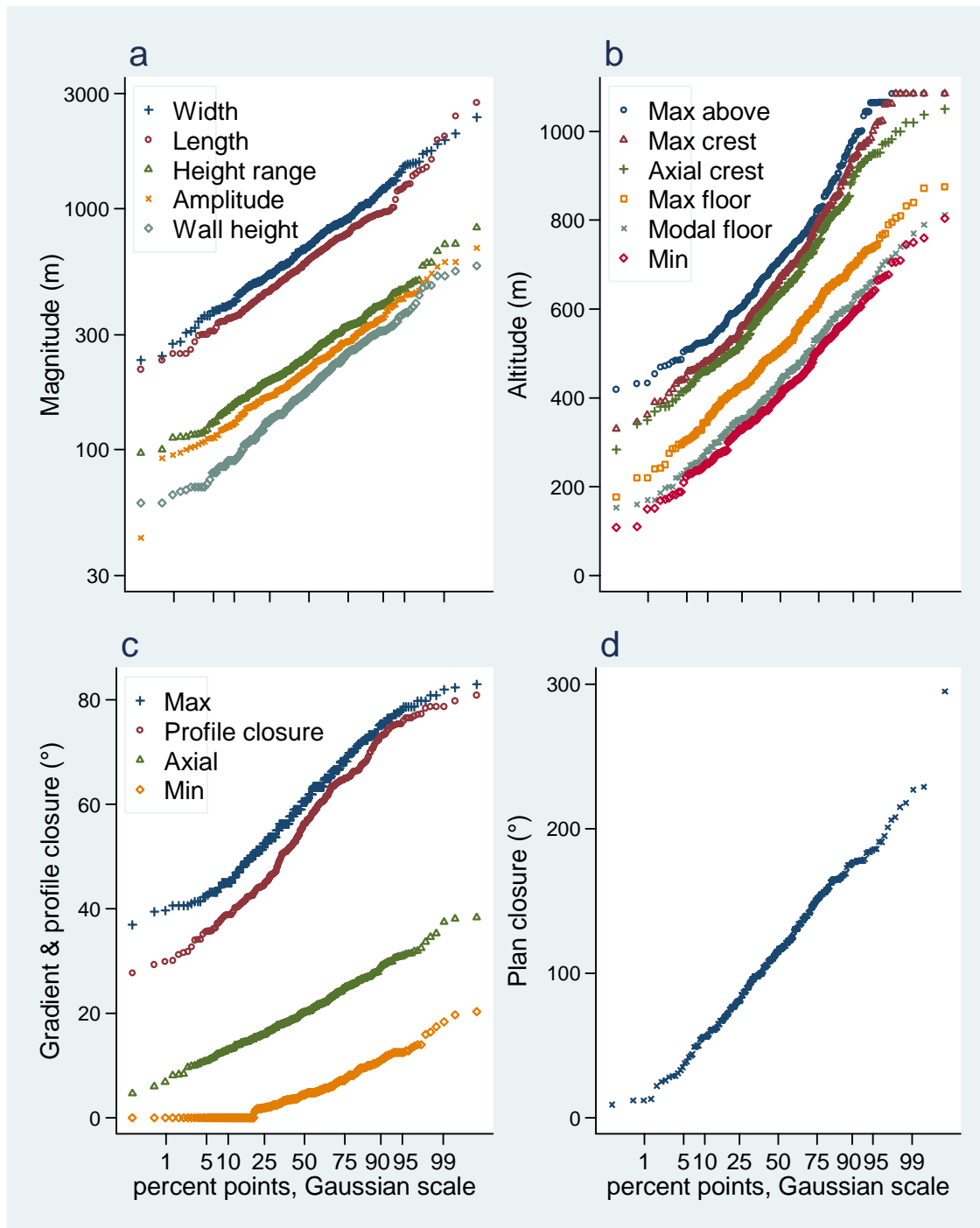


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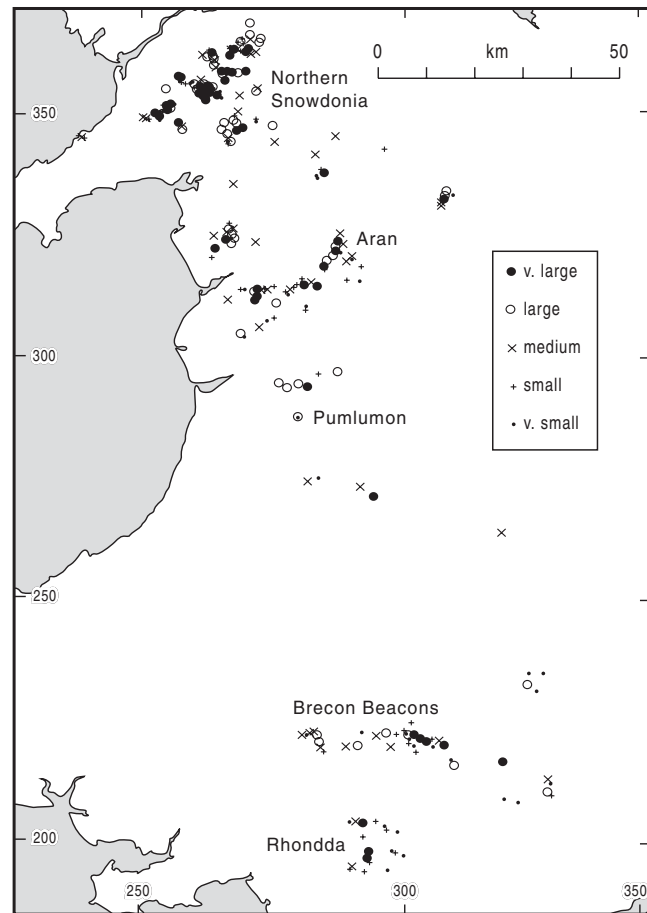


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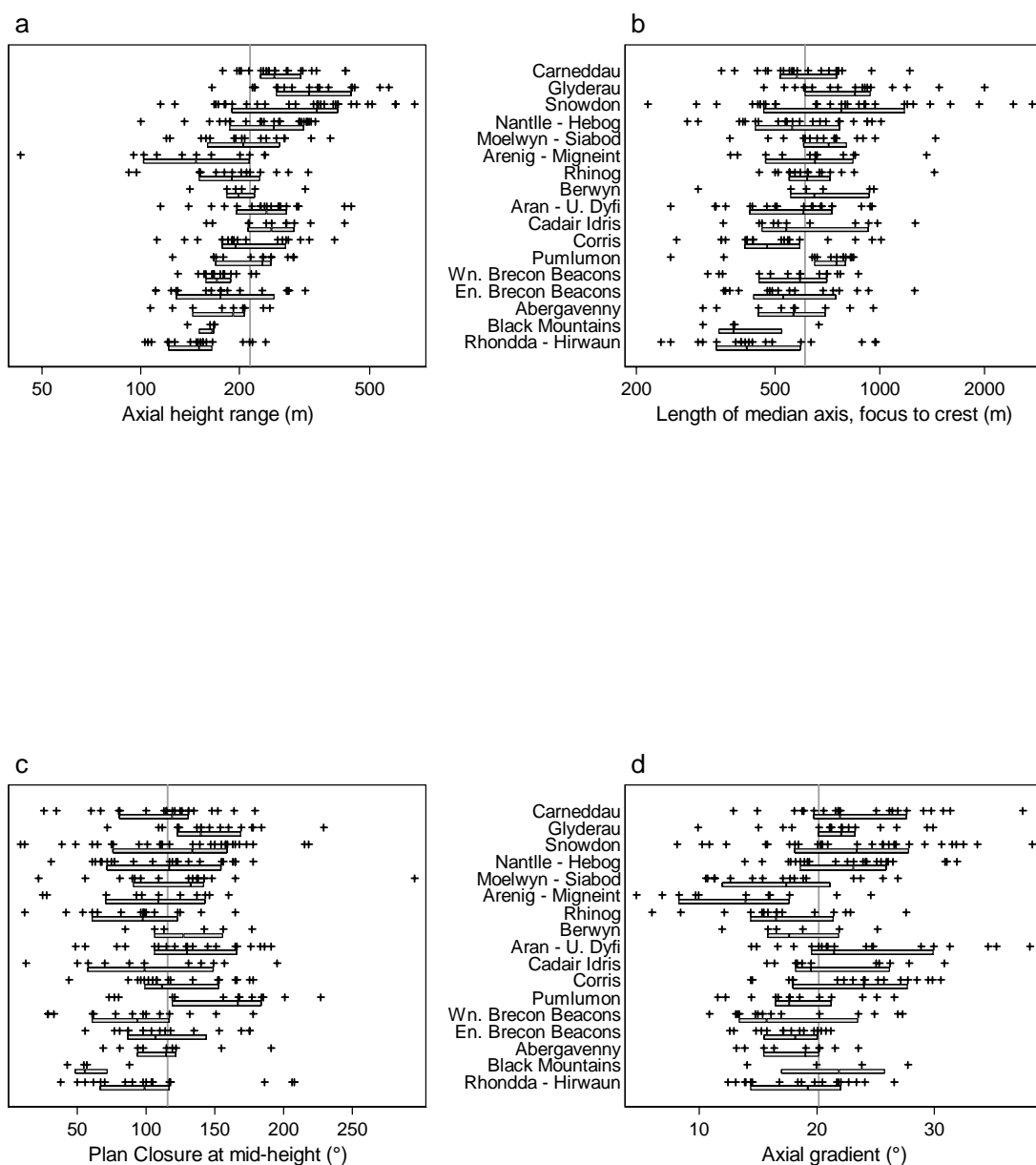
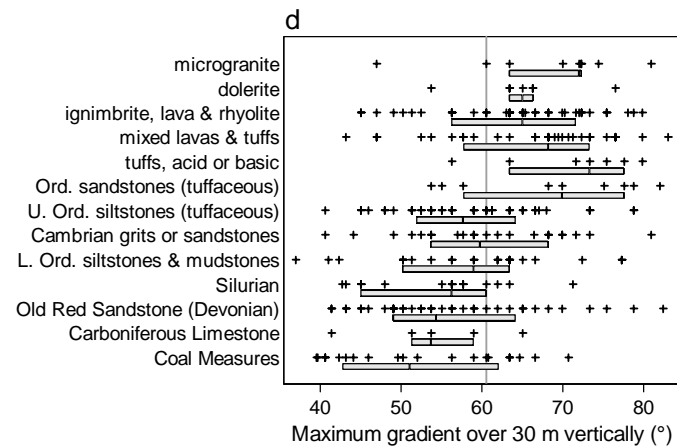
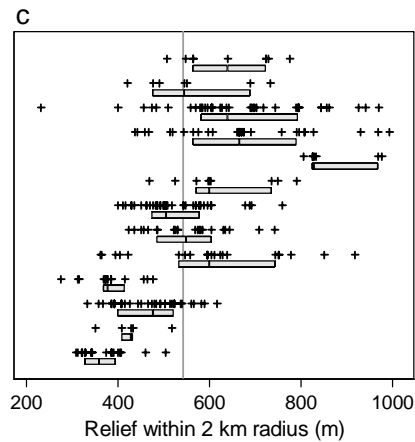
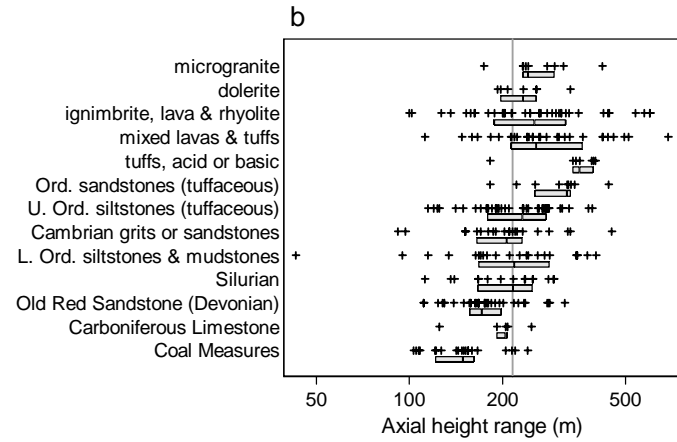
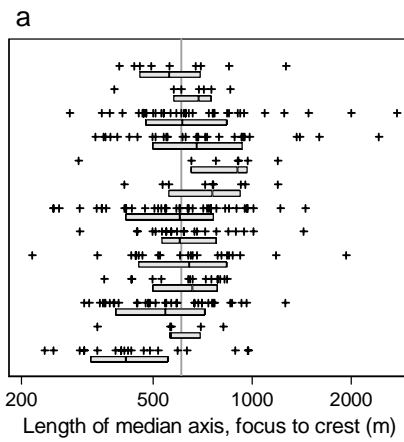


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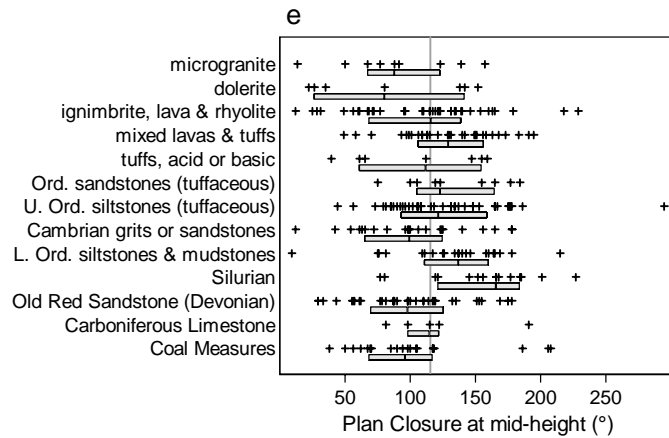


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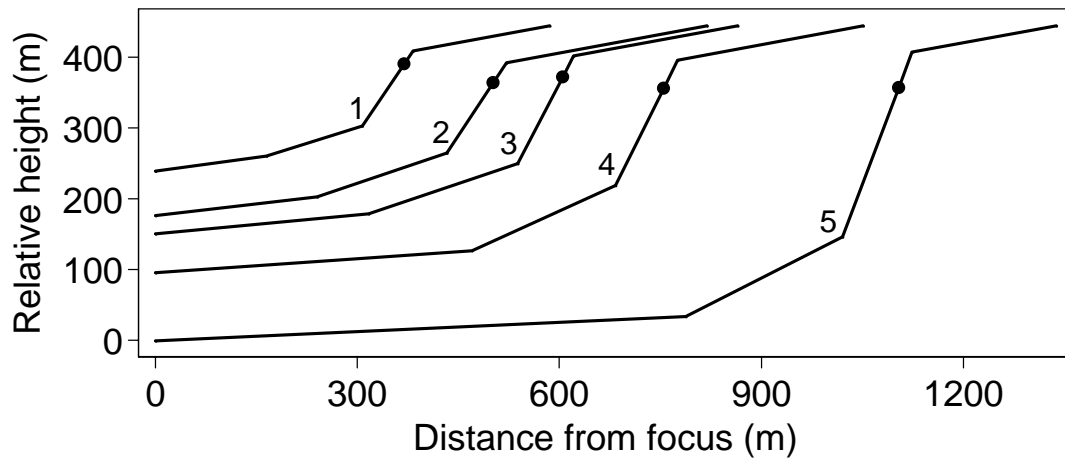


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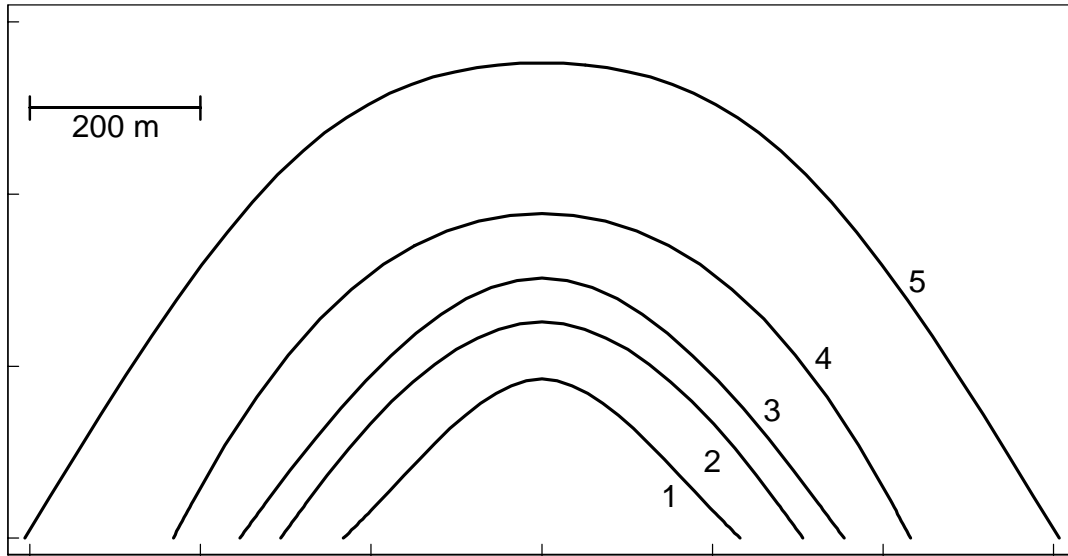


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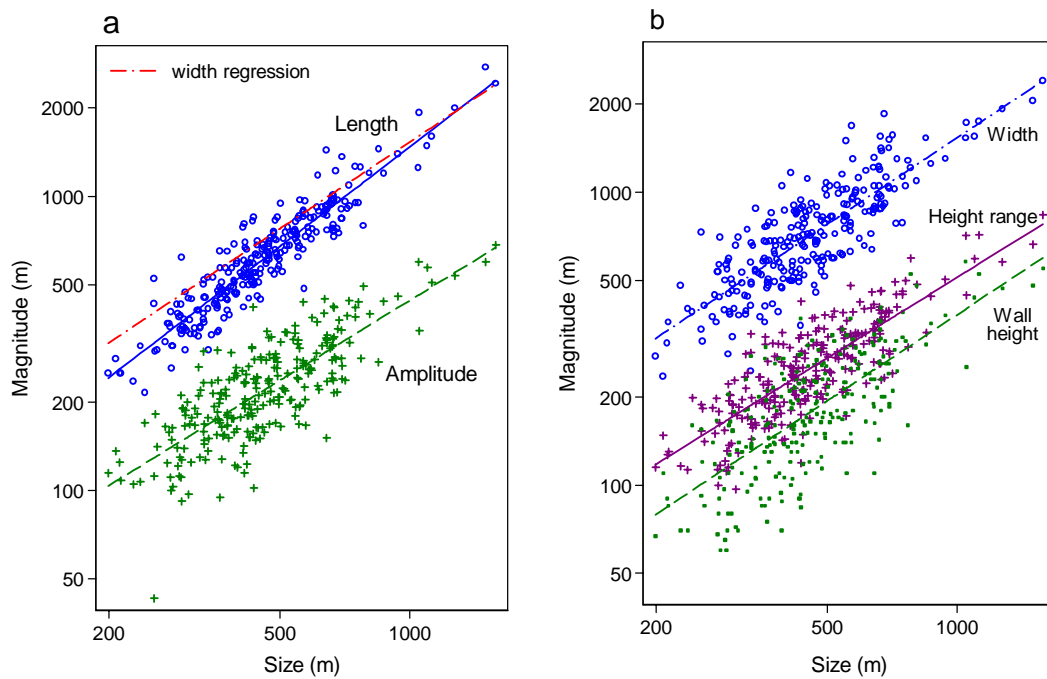


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Dear Takashi,

I have made the necessary figure corrections and just some tiny further changes to the text - removing one more sentence (after Davis 1911 sentence) in 1.2. If you wish I can send the corresponding text with changes.

Many thanks for your help,

Ian.



|     |      |      |     |     |     |     |     |     |      |      |     |      |      |      |     |                |          |              |            |                |             |            |            |     |              |          |          |          |          |          |          |          |          |           |          |          |          |          |          |   |
|-----|------|------|-----|-----|-----|-----|-----|-----|------|------|-----|------|------|------|-----|----------------|----------|--------------|------------|----------------|-------------|------------|------------|-----|--------------|----------|----------|----------|----------|----------|----------|----------|----------|-----------|----------|----------|----------|----------|----------|---|
| 116 | 2875 | 3254 | 439 | 445 | 485 | 710 | 720 | 862 | 683  | 520  | 225 | 72.3 | 0    | 48   | 522 | 607 major mort | 0        | well-defne   | LLS        | glacie         | valley-side | Arari-Dyfi | mixed lava | 84  | 82           | Llwynbar | 243      | 48       | 19.58471 | 723      | 0.761347 | 8.63E+07 | 441.9198 | 21.283421 | 2.716003 | 2.352183 | 2.654354 | 1.602758 | 2.432989 | 0 |
| 117 | 2870 | 3260 | 385 | 430 | 590 | 625 | 684 | 590 | 624  | 590  | 624 | 73.1 | 0    | 49   | 525 | 607 major mort | 0        | well-defne   | LLS        | glacie         | valley-side | Arari-Dyfi | mixed lava | 86  | 81           | Cwmlyf   | 439      | 205      | 24.04458 | 3.021169 | 4.249138 | 2.6806   | 2.432989 | 2.352183  | 2.654354 | 1.602758 | 2.432989 | 0        |          |   |
| 118 | 2866 | 3227 | 577 | 577 | 725 | 685 | 685 | 685 | 685  | 685  | 685 | 68.9 | 0    | 129  | 462 | 575 major mort | 0        | well-defne   | uncertain  | valley-side    | Arari-Dyfi  | mixed lava | 135        | 68  | Arari-Dyfi   | 265      | 148      | 16.69561 | 69.0     | 0.887797 | 1.80E+08 | 584.7421 | 2.432989 | 2.352183  | 2.654354 | 1.602758 | 2.432989 | 0        |          |   |
| 119 | 2866 | 3217 | 500 | 500 | 670 | 901 | 905 | 754 | 904  | 1217 | 312 | 69   | 1.7  | 151  | 535 | 669 diff       | 0        | define       | uncertain  | valley-side    | Arari-Dyfi  | mixed lava | 51         | 95  | Llithnam     | 148      | 184      | 18.83437 | 67.3     | 1.294861 | 2.86E+08 | 658.8494 | 2.432989 | 2.352183  | 2.654354 | 1.602758 | 2.432989 | 0        |          |   |
| 120 | 2861 | 3232 | 360 | 385 | 495 | 590 | 590 | 590 | 590  | 590  | 590 | 60.5 | 1.3  | 176  | 360 | 587 diff       | 0        | 2 poor       | uncertain  | valley-side    | Arari-Dyfi  | mixed lava | 117        | 147 | Cwmlyf       | 116      | 117      | 14.74444 | 57.2     | 1.289953 | 2.49E+08 | 2.49E+08 | 2.432989 | 2.352183  | 2.654354 | 1.602758 | 2.432989 | 0        |          |   |
| 121 | 2871 | 3214 | 605 | 605 | 720 | 720 | 720 | 720 | 720  | 720  | 720 | 72.5 | 0    | 158  | 570 | 570 diff       | 0        | 1 poor       | uncertain  | valley-side    | Arari-Dyfi  | mixed lava | 115        | 50  | Arari-Dyfi   | 115      | 50       | 24.70024 | 34.1     | 1.1      | 7.900025 | 1.69E+07 | 2.432989 | 2.352183  | 2.654354 | 1.602758 | 2.432989 | 0        |          |   |
| 122 | 2861 | 3208 | 380 | 470 | 555 | 770 | 770 | 648 | 770  | 648  | 770 | 64.8 | 10   | 134  | 589 | 758 diff       | 0        | 1 define     | uncertain  | valley-side    | Arari-Dyfi  | mixed lava | 120        | 115 | Gwyn-y-llw   | 268      | 175      | 20.15936 | 55       | 1.310588 | 1.88E+08 | 572.6769 | 2.432989 | 2.352183  | 2.654354 | 1.602758 | 2.432989 | 0        |          |   |
| 123 | 2867 | 3196 | 298 | 280 | 344 | 504 | 732 | 497 | 642  | 675  | 190 | 59   | 4.8  | 186  | 381 | 543 outcrops   | 0        | 1 well-defne | uncertain  | valley-side    | Arari-Dyfi  | mixed lava | 119        | 117 | Dwyn-y-llw   | 239      | 86       | 24.11409 | 54.2     | 1.051402 | 1.04E+08 | 572.6769 | 2.432989 | 2.352183  | 2.654354 | 1.602758 | 2.432989 | 0        |          |   |
| 124 | 2868 | 3157 | 182 | 230 | 438 | 432 | 414 | 608 | 498  | 432  | 414 | 60.8 | 10.7 | 106  | 329 | 498 diff       | 0        | 1 marginal   | none       | valley-side    | Arari-Dyfi  | mixed lava | 130        | 128 | Cwmlyf       | 232      | 128      | 20.84762 | 29.0     | 0.821782 | 7.05E+07 | 412.1376 | 2.432989 | 2.352183  | 2.654354 | 1.602758 | 2.432989 | 0        |          |   |
| 125 | 2916 | 3185 | 209 | 220 | 286 | 463 | 598 | 408 | 500  | 770  | 193 | 57.7 | 10.7 | 135  | 438 | 464 diff       | 0        | 1 poor       | uncertain  | valley-side    | Arari-Dyfi  | mixed lava | 130        | 350 | Penn-y-gelli | 197      | 67       | 17.03044 | 47       | 1.6      | 6.90E+07 | 410.0575 | 2.432989 | 2.352183  | 2.654354 | 1.602758 | 2.432989 | 0        |          |   |
| 126 | 2911 | 3155 | 330 | 340 | 390 | 484 | 525 | 470 | 430  | 396  | 97  | 48   | 5.4  | 166  | 230 | 375 diff       | 0        | 0 poor       | uncertain  | valley-side    | Arari-Dyfi  | mixed lava | 346        | 355 | Llan-y-fawr  | 140      | 77       | 18.04249 | 42.6     | 0.92023  | 1.98E+07 | 287.043  | 2.432989 | 2.352183  | 2.654354 | 1.602758 | 2.432989 | 0        |          |   |
| 127 | 2948 | 3198 | 236 | 265 | 380 | 600 | 879 | 540 | 600  | 879  | 540 | 73.5 | 0    | 183  | 570 | 660 diff       | 0        | 2 define     | none       | valley-side    | Arari-Dyfi  | mixed lava | 141        | 80  | Cwmlyf       | 304      | 144      | 24.18108 | 66.5     | 1.21811  | 1.70E+08 | 653.7394 | 2.432989 | 2.352183  | 2.654354 | 1.602758 | 2.432989 | 0        |          |   |
| 128 | 2845 | 3190 | 350 | 385 | 400 | 652 | 780 | 650 | 424  | 360  | 270 | 61.9 | 19.7 | 115  | 646 | 664 outcrops   | 0        | 0 poor       | uncertain  | valley-side    | Arari-Dyfi  | mixed lava | 140        | 130 | N Cwyrdd     | 300      | 50       | 35.28121 | 42.2     | 0.84057  | 4.58E+07 | 357.7639 | 2.432989 | 2.352183  | 2.654354 | 1.602758 | 2.432989 | 0        |          |   |
| 129 | 2841 | 3186 | 420 | 465 | 480 | 485 | 700 | 754 | 683  | 335  | 363 | 230  | 68.2 | 10.3 | 130 | 663 outcrops   | 0        | 0 marginal   | uncertain  | inner          | cont        | Arari-Dyfi | mixed lava | 88  | 92           | L Cwyrdd | 263      | 65       | 38.13456 | 47.0     | 1.432384 | 3.37E+07 | 327.1467 | 2.432989  | 2.352183 | 2.654354 | 1.602758 | 2.432989 | 0        |   |
| 130 | 2751 | 3110 | 253 | 290 | 455 | 534 | 545 | 524 | 852  | 787  | 150 | 51.3 | 6    | 134  | 419 | 451 diff       | 0        | 1 poor       | none       | valley-side    | Arari-Dyfi  | mixed lava | 81         | 30  | Bychan       | 271      | 202      | 17.75547 | 45.3     | 0.923709 | 1.82E+08 | 566.498  | 2.432989 | 2.352183  | 2.654354 | 1.602758 | 2.432989 | 0        |          |   |
| 131 | 2845 | 3179 | 350 | 385 | 400 | 637 | 584 | 602 | 583  | 338  | 528 | 200  | 52.5 | 18.4 | 79  | 570            | 678 diff | 0            | 0 marginal | none           | valley-side | Arari-Dyfi | mixed lava | 50  | 33           | Yhen     | 233      | 87       | 34.5804  | 34.1     | 1.56213  | 4.16E+07 | 346.446  | 2.432989  | 2.352183 | 2.654354 | 1.602758 | 2.432989 | 0        |   |
| 132 | 2830 | 3144 | 175 | 200 | 330 | 434 | 674 | 655 | 950  | 1130 | 390 | 66.6 | 5    | 153  | 515 | 570 diff       | 0        | 0 define     | uncertain  | valley-side    | Arari-Dyfi  | mixed lava | 97         | 60  | Thames       | 300      | 125      | 22.21934 | 61.6     | 1.189474 | 4.19E+08 | 748.093  | 2.432989 | 2.352183  | 2.654354 | 1.602758 | 2.432989 | 0        |          |   |
| 133 | 2855 | 3147 | 296 | 360 | 420 | 595 | 659 | 594 | 1010 | 875  | 260 | 63.4 | 5    | 178  | 425 | 595 diff       | 0        | 3 well-defne | none       | valley-side    | Arari-Dyfi  | mixed lava | 24         | 45  | Porta-Cru    | 328      | 154      | 17.99133 | 58.0     | 0.866337 | 2.95E+08 | 681.8117 | 2.432989 | 2.352183  | 2.654354 | 1.602758 | 2.432989 | 0        |          |   |
| 134 | 2791 | 3148 | 432 | 470 | 560 | 654 | 664 | 610 | 550  | 675  | 135 | 55   | 12.5 | 175  | 378 | 463 diff       | 0        | 1 poor       | LLS        | snopwally-side | Arari-Dyfi  | mixed lava | 134        | 142 | W. Wain-o    | 178      | 118      | 12.73347 | 42.5     | 1.222723 | 6.61E+07 | 404.2923 | 2.432989 | 2.352183  | 2.654354 | 1.602758 | 2.432989 | 0        |          |   |
| 135 | 2785 | 3142 | 443 | 500 | 570 | 636 | 670 | 636 | 432  | 522  | 85  | 48   | 10.3 | 94   | 430 | 482 diff       | 0        | 0 marginal   | uncertain  | valley-side    | Arari-Dyfi  | mixed lava | 193        | 127 | 0.74315      | 37       | 1.206333 | 4.35E+07 | 351.7525 | 2.432989 | 2.352183 | 2.654354 | 1.602758 | 2.432989  | 0        |          |          |          |          |   |
| 136 | 2778 | 3138 | 310 | 440 | 500 | 610 | 625 | 572 | 544  | 800  | 114 | 59.3 | 17.4 | 165  | 448 | 492 diff       | 0        | 0 marginal   | none       | valley-side    | Arari-Dyfi  | mixed lava | 262        | 129 | 20.71629     | 38.9     | 1.470588 | 1.14E+08 | 684.9125 | 2.432989 | 2.352183 | 2.654354 | 1.602758 | 2.432989  | 0        |          |          |          |          |   |
| 137 | 2769 | 3133 | 370 | 430 | 490 | 590 | 590 | 555 | 416  | 593  | 85  | 48   | 10   | 108  | 395 | 476 diff       | 0        | 0 marginal   | none       | valley-side    | Arari-Dyfi  | mixed lava | 120        | 95  | N Yhen       | 185      | 120      | 23.97527 | 38       | 1.425481 | 2.19E+07 | 357.3605 | 2.432989 | 2.352183  | 2.654354 | 1.602758 | 2.432989 | 0        |          |   |
| 138 | 2772 | 3127 | 330 | 355 | 430 | 508 | 508 | 479 | 280  | 596  | 90  | 53.7 | 9.6  | 44   | 402 | 485 diff       | 0        | 0 poor       | none       | valley-side    | Arari-Dyfi  | mixed lava | 43         | 55  | N Yhen       | 149      | 100      | 29.81903 | 44.1     | 2.176623 | 2.19E+07 | 279.893  | 2.432989 | 2.352183  | 2.654354 | 1.602758 | 2.432989 | 0        |          |   |
| 139 | 2751 | 3110 | 253 | 290 | 455 | 534 | 545 | 524 | 852  | 787  | 150 | 51.3 | 6    | 134  | 419 | 451 diff       | 0        | 1 poor       | none       | valley-side    | Arari-Dyfi  | mixed lava | 81         | 30  | Bychan       | 271      | 202      | 17.75547 | 45.3     | 0.923709 | 1.82E+08 | 566.498  | 2.432989 | 2.352183  | 2.654354 | 1.602758 | 2.432989 | 0        |          |   |
| 140 | 2747 | 3079 | 152 | 170 | 220 | 362 | 486 | 350 | 350  | 675  | 160 | 61.2 | 7.8  | 87   | 416 | 542 diff       | 0        | 0 marginal   | none       | valley-side    | Arari-Dyfi  | mixed lava | 58         | 46  | Cae-coch     | 198      | 68       | 29.44768 | 53.6     | 1.925741 | 4.68E+07 | 390.3122 | 2.432989 | 2.352183  | 2.654354 | 1.602758 | 2.432989 | 0        |          |   |
| 141 | 2718 | 3060 | 347 | 370 | 460 | 647 | 658 | 625 | 590  | 610  | 227 | 56.3 | 9.6  | 116  | 422 | 504 diff       | 0        | 0 poor       | none       | valley-side    | Arari-Dyfi  | mixed lava | 86         | 80  | Drifd Tfar   | 178      | 115      | 28.29218 | 47       | 1.03398  | 1.00E+08 | 464.2396 | 2.432989 | 2.352183  | 2.654354 | 1.602758 | 2.432989 | 0        |          |   |
| 142 | 2887 | 3040 | 418 | 450 | 495 | 622 | 634 | 608 | 550  | 390  | 127 | 49.1 | 13.4 | 105  | 423 | 564 diff       | 0        | 0 poor       | none       | valley-side    | Arari-Dyfi  | mixed lava | 95         | 81  | Tarrent      | 190      | 77       | 28.40654 | 35.7     | 1.14228  | 2.50E+07 | 260.025  | 2.432989 | 2.352183  | 2.654354 | 1.602758 | 2.432989 | 0        |          |   |
| 143 | 2852 | 3047 | 307 | 330 | 400 | 607 | 634 | 590 | 770  | 585  | 77  | 52.5 | 6    | 96   | 434 | 576 diff       | 0        | 0 define     | none       | valley-side    | Arari-Dyfi  | mixed lava | 359        | 28  | N Tarrent    | 83       | 123      | 21.73181 | 46.5     | 0.823944 | 1.18E+08 | 489.8542 | 2.432989 | 2.352183  | 2.654354 | 1.602758 | 2.432989 | 0        |          |   |
| 144 | 2857 | 3117 | 351 | 354 | 400 | 584 | 589 | 510 | 540  | 860  | 215 | 53.7 | 0    | 140  | 337 | 514 major mort | 0        | 0 well-defne | uncertain  | valley-side    | Arari-Dyfi  | mixed lava | 6          | 343 | L Cyn        | 159      | 49       | 16.40877 | 53.7     | 1.592939 | 7.38E+07 | 2.201397 | 2.432989 | 2.352183  | 2.654354 | 1.602758 | 2.432989 | 0        |          |   |
| 145 | 2862 | 3138 | 390 | 440 | 495 | 610 | 643 | 603 | 445  | 590  | 163 | 63.3 | 6    | 58   | 506 | 660 diff       | 0        | 0 define     | uncertain  | valley-side    | Arari-Dyfi  | mixed lava | 14         | 35  | L Cyn        | 133      | 102      | 25.33149 | 57.4     | 1.22222  | 4.94E+07 | 464.415  | 2.432989 | 2.352183  | 2.654354 | 1.602758 | 2.432989 | 0        |          |   |
| 146 | 2868 | 3135 | 405 | 415 | 465 | 572 | 616 | 588 | 585  | 680  | 115 | 69   | 11   | 99   | 519 | 620 diff       | 0        | 0 marginal   | uncertain  | valley-side    | Arari-Dyfi  | mixed lava | 115        | 69  | 21.071 Carr  | 208      | 145      | 25.06778 | 27.7     | 1.02778  | 2.69E+08 | 2.432989 | 2.352183 | 2.654354  | 1.602758 | 2.432989 | 0        |          |          |   |
| 147 | 2709 | 3115 | 422 | 480 | 583 | 763 | 791 | 752 | 930  | 850  | 190 | 57.7 | 4.8  | 131  | 593 | 808 diff       | 0        | 0 define     | LLS        | snopwally-side | Arari-Dyfi  | mixed lava | 152        | 140 | Amarrh       | 330      | 101      | 26.50668 | 52.9     | 0.913979 | 2.61E+08 | 638.9575 | 2.432989 | 2.352183  | 2.654354 | 1.602758 | 2.432989 | 0        |          |   |
| 148 | 2714 | 3124 | 445 | 470 | 568 | 650 | 683 | 723 | 985  | 1050 | 365 | 76.5 | 0    | 195  | 518 | 806 major rock | 0        | 1 classic    | LLS        | glacie         | valley-side | Arari-Dyfi | mixed lava | 85  | 65           | L Cyn    | 278      | 123      | 15.76094 | 76.5     | 1.06599  | 2.98E+08 | 680.0195 | 2.432989  | 2.352183 | 2.654354 | 1.602758 | 2.432989 | 0        |   |
| 149 | 2714 | 3124 | 445 | 470 | 568 | 650 | 683 | 723 | 985  | 1050 | 365 | 76.5 | 0    | 195  | 518 | 806 major rock | 0        | 1 classic    | LLS        | glacie         | valley-side | Arari-Dyfi | mixed lava | 85  | 65           | L Cyn    | 278      | 123      | 15.76094 | 76.5     | 1.06599  | 2.98E+08 | 680.0195 | 2.432989  | 2.352183 | 2.654354 | 1.602758 | 2.432989 | 0        |   |
| 150 | 2714 | 3138 | 403 | 412 | 575 | 852 | 852 | 823 | 1265 | 784  | 320 | 72.2 | 0    | 50   | 525 | 723 major mort | 0        | 1 define     | LLS        | snopwally-side | Arari-Dyfi  |            |            |     |              |          |          |          |          |          |          |          |          |           |          |          |          |          |          |   |



|     |      |      |     |     |     |     |     |     |     |      |     |      |      |     |     |                 |              |  |     |                 |     |     |          |      |          |          |          |     |          |          |          |          |          |          |          |          |   |
|-----|------|------|-----|-----|-----|-----|-----|-----|-----|------|-----|------|------|-----|-----|-----------------|--------------|--|-----|-----------------|-----|-----|----------|------|----------|----------|----------|-----|----------|----------|----------|----------|----------|----------|----------|----------|---|
| 241 | 2919 | 2035 | 330 | 365 | 440 | 559 | 600 | 545 | 975 | 1408 | 180 | 64.6 | 0    | 100 | 319 | 385 major mori  | 0 classic    | LLS glacie valley-side Rhondaa+Coal Meas     | 37  | 31 L. Fawr, G   | 215 | 110 | 12.43544 | 64.6 | 1.444103 | 2.95E+08 | 865.8073 | 229 | 2.989005 | 3.148803 | 2.256272 | 2.332438 | 2.823349 | 2.041933 | 2.359835 | 0        |   |
| 242 | 2944 | 2038 | 370 | 399 | 438 | 512 | 515 | 492 | 470 | 1050 | 88  | 60.9 | 2.7  | 118 | 255 | 320 major bog   | 0 poor       | LLS glacie valley-side Rhondaa+Coal Meas     | 45  | 38 C-y-Bwch     | 122 | 68  | 14.55138 | 58.3 | 2.234043 | 6.02E+07 | 291.5365 | 142 | 2.672098 | 2.921189 | 1.944463 | 2.08638  | 2.583216 | 1.632509 | 2.162288 | 0        |   |
| 243 | 2959 | 2028 | 358 | 400 | 440 | 475 | 510 | 461 | 310 | 700  | 80  | 38.7 | 3.4  | 38  | 256 | 329 drift       | 0 marginal   | uncertain valley-side Rhondaa+Coal Meas      | 62  | 47 C-y-yr-Ysgo  | 103 | 82  | 18.37948 | 36.3 | 2.256065 | 2.24E+07 | 281.6862 | 117 | 2.491362 | 2.845098 | 1.778151 | 2.012837 | 2.449786 | 1.913814 | 2.068186 | 0        |   |
| 244 | 2965 | 2020 | 237 | 248 | 312 | 393 | 510 | 392 | 430 | 680  | 130 | 42.3 | 3.7  | 98  | 257 | 342 drift       | 0 poor       | uncertain valley-hearRhondaa+Coal Meas       | 24  | 18 Tarren-y-B   | 155 | 75  | 19.82246 | 38.6 | 1.581395 | 4.53E+07 | 356.5357 | 156 | 2.633468 | 2.832509 | 2.113943 | 2.190332 | 2.552103 | 1.875081 | 2.193125 | 0        |   |
| 245 | 2984 | 2016 | 275 | 290 | 346 | 391 | 433 | 380 | 235 | 480  | 70  | 44.1 | 10.1 | 86  | 275 | 308 drift       | 0 marginal   | uncertain valley-side Rhondaa+Coal Meas      | 0   | 385 Cefnhoec-r  | 105 | 71  | 24.0755  | 34   | 2.042553 | 1.18E+07 | 227.9464 | 116 | 2.371988 | 2.681241 | 1.845098 | 2.021189 | 2.387833 | 1.851258 | 2.064458 | 0        |   |
| 246 | 2919 | 2008 | 330 | 365 | 396 | 490 | 522 | 480 | 300 | 880  | 105 | 56.3 | 5.7  | 62  | 309 | 343 drift       | 0 poor       | uncertain valley-side Rhondaa+Coal Meas      | 100 | 97 C. Blaenrh   | 150 | 66  | 26.66505 | 50.6 | 2.933333 | 3.96E+07 | 340.8514 | 160 | 2.477121 | 2.984483 | 2.021189 | 2.176991 | 2.532556 | 1.815644 | 2.20412  | 0        |   |
| 247 | 2930 | 1978 | 253 | 330 | 417 | 514 | 523 | 494 | 970 | 1224 | 143 | 60.5 | 4.4  | 206 | 333 | 386 drift       | 0 well-defne | LLS glacie valley-side Rhondaa+Coal Meas     | 44  | 55 Saerhren,    | 241 | 164 | 13.95282 | 56.1 | 1.261656 | 2.86E+08 | 658.9565 | 261 | 2.986772 | 3.087781 | 2.155336 | 2.382017 | 2.818857 | 2.214844 | 2.416641 | 0        |   |
| 248 | 2927 | 1962 | 321 | 350 | 412 | 544 | 559 | 541 | 890 | 1184 | 165 | 66.6 | 2    | 208 | 308 | 373 major bog   | 0 well-defne | LLS glacie valley-hear Rhondaa+Coal Meas     | 80  | 76 Ffarc, Cwfr  | 220 | 91  | 13.88467 | 64.6 | 1.330337 | 2.32E+08 | 614.3108 | 223 | 2.949309 | 3.073352 | 2.217484 | 2.324243 | 2.788388 | 1.969041 | 2.248305 | 0        |   |
| 249 | 2932 | 1953 | 337 | 350 | 410 | 508 | 524 | 503 | 410 | 785  | 146 | 63.4 | 3.1  | 104 | 337 | 388 major bog   | 0 well-defne | LLS glacie valley-side Rhondaa+Coal Meas     | 43  | 48 Crig Ffard   | 166 | 73  | 22.04194 | 60.3 | 1.914634 | 5.34E+07 | 378.6349 | 171 | 2.818784 | 2.89487  | 2.164353 | 2.220108 | 2.579521 | 1.863323 | 2.232966 | 0        |   |
| 250 | 2898 | 1945 | 308 | 320 | 427 | 537 | 565 | 460 | 594 | 854  | 130 | 49.6 | 4.4  | 186 | 329 | 399 drift       | 1 well-defne | LLS glacie valley-hearRhondaa+Coal Meas      | 146 | 128 Bluenagw    | 152 | 119 | 14.35354 | 45.2 | 1.437771 | 7.71E+07 | 425.6271 | 229 | 2.773787 | 2.931458 | 2.113943 | 2.181844 | 2.629029 | 2.075547 | 2.359835 | 1        |   |
| 251 | 2965 | 1937 | 330 | 350 | 378 | 460 | 472 | 438 | 250 | 360  | 85  | 40.6 | 11.3 | 105 | 259 | 390 drift       | 0 marginal   | none valley-side Rhondaa+Coal Meas           | 33  | 21 Bwlffu, Dwr  | 108 | 48  | 23.36434 | 29.3 | 1.44     | 9720000  | 213.4136 | 130 | 2.39794  | 2.556303 | 1.929419 | 2.033424 | 2.339222 | 1.681241 | 2.113943 | 0        |   |
| 252 | 2973 | 1977 | 350 | 365 | 413 | 471 | 481 | 471 | 400 | 500  | 65  | 46   | 4.8  | 70  | 229 | 328 drift       | 0 marginal   | uncertain valley-side Rhondaa+Coal Meas      | 29  | 40 Maerdy, C    | 121 | 63  | 16.83057 | 41.2 | 1.25     | 2.43E+07 | 289.249  | 121 | 2.60206  | 2.68697  | 1.812913 | 2.082785 | 2.461272 | 1.796341 | 2.082785 | 0        |   |
| 253 | 2982 | 1973 | 282 | 302 | 364 | 462 | 475 | 424 | 420 | 1165 | 100 | 59   | 2.4  | 67  | 252 | 321 drift       | 0 marginal   | uncertain valley-side Rhondaa+Coal Meas      | 25  | 15 Tarren Ma    | 142 | 82  | 18.68015 | 56.6 | 2.773809 | 6.95E+07 | 411.1067 | 180 | 2.623249 | 3.066326 |          | 2        | 2.152288 | 2.613955 | 1.913814 | 2.255272 | 0 |
| 254 | 2996 | 1967 | 253 | 254 | 315 | 390 | 419 | 380 | 340 | 593  | 110 | 63.4 | 0    | 50  | 231 | 329 major mori  | 0 poor       | uncertain valley-side Rhondaa+Coal Meas      | 73  | 77 C. Rhond     | 127 | 62  | 20.48214 | 63.4 | 1.744118 | 2.56E+07 | 294.7445 | 137 | 2.531470 | 2.773055 | 2.041393 | 2.103804 | 2.469444 | 1.762392 | 2.136721 | 0        |   |
| 255 | 2984 | 1939 | 320 | 359 | 396 | 530 | 555 | 525 | 490 | 471  | 135 | 39.5 | 8.3  | 117 | 367 | 407 drift       | 0 marginal   | uncertain valley-side Rhondaa+Coal Meas      | 145 | 145 Dwyw, Cw    | 205 | 78  | 22.70284 | 31.2 | 0.961225 | 4.73E+07 | 361.6793 | 210 | 2.690186 | 2.873021 | 2.130334 | 2.311754 | 2.558324 | 1.880814 | 2.322219 | 0        |   |
| 256 | 2922 | 1934 | 370 | 390 | 434 | 532 | 537 | 525 | 386 | 850  | 130 | 50.2 | 3.7  | 85  | 339 | 404 drift       | 0 marginal   | uncertain valley-side Rhondaa+Coal Meas      | 73  | 80 Ffurch, Tai  | 155 | 64  | 21.87815 | 46.5 | 2.202073 | 5.09E+07 | 370.4924 | 162 | 2.586587 | 2.929419 | 2.113943 | 2.190332 | 2.568779 | 1.80618  | 2.209515 | 0        |   |
| 270 | 2368 | 3456 | 110 | 153 | 177 | 330 | 564 | 284 | 440 | 580  | 170 | 80.9 | 5.7  | 139 | 541 | 564 drift       | 0 definite   | uncertain valley-side Nantlle-He microgranit | 34  | 25 Llwyd-y-gr   | 174 | 67  | 21.57655 | 75.2 | 1.318182 | 4.44E+07 | 354.1142 | 220 | 2.643453 | 2.763428 | 2.230449 | 2.240549 | 2.549143 | 1.826075 | 2.342423 | 0        |   |
| 271 | 2374 | 3453 | 150 | 170 | 220 | 410 | 564 | 382 | 675 | 684  | 207 | 63.4 | 3.8  | 91  | 512 | 564 drift       | 0 definite   | uncertain valley-side Nantlle-He microgranit | 16  | 23 Ceiliog, G   | 232 | 70  | 18.98003 | 59.6 | 0.983704 | 1.04E+08 | 470.2404 | 260 | 2.825304 | 2.822168 | 2.31597  | 2.365488 | 2.67322  | 1.845098 | 2.414873 | 0        |   |
| 272 | 2379 | 3451 | 108 | 160 | 240 | 345 | 485 | 342 | 562 | 538  | 140 | 47   | 8.1  | 77  | 431 | 547 drift       | 0 marginal   | uncertain valley-side Nantlle-He microgranit | 69  | 60 Merbyll, Ti  | 234 | 132 | 22.60538 | 38.9 | 0.953737 | 7.05E+07 | 413.0846 | 237 | 2.749738 | 2.729165 | 2.146128 | 2.368216 | 2.616039 | 2.120574 | 2.374748 | 0        |   |
| 273 | 2718 | 3649 | 372 | 372 | 405 | 695 | 790 | 630 | 754 | 1015 | 295 | 66.3 | 0    | 80  | 384 | 477 major mori  | 0 poor       | LLS glacie valley-side Carneddau dolerite    | 105 | 95 S. C. Eglau  | 258 | 33  | 18.88967 | 66.3 | 1.346154 | 1.97E+08 | 882.3075 | 323 | 2.877371 | 3.008466 | 2.469822 | 2.41162  | 2.765152 | 1.518514 | 2.509202 | 0        |   |
| 274 | 2721 | 3657 | 372 | 372 | 395 | 631 | 729 | 629 | 715 | 1120 | 250 | 66.3 | 0    | 26  | 390 | 420 major mori  | 0 marginal   | LLS glacie valley-side Carneddau dolerite    | 106 | 106 N. C. Eglau | 257 | 22  | 19.77049 | 66.3 | 1.568434 | 2.06E+08 | 590.6822 | 259 | 2.854306 | 3.040219 | 2.4133   | 2.409833 | 2.771152 | 1.961728 | 2.4133   | 0        |   |
| 281 | 2643 | 3534 | 431 | 460 | 465 | 614 | 619 | 613 | 297 | 310  | 120 | 56.3 | 15.9 | 61  | 628 | 828 outcropping | 0 marginal   | uncertain valley-side Snowdon luffs, acid r  | 28  | 35 Wenallt, G   | 182 | 64  | 31.49974 | 40.4 | 1.043771 | 1.68E+07 | 255.8958 | 183 | 2.472756 | 2.491382 | 2.079181 | 2.260071 | 2.408063 | 1.80618  | 2.282451 | 0        |   |
| 283 | 2628 | 3567 | 396 | 420 | 507 | 675 | 921 | 635 | 660 | 625  | 370 | 70.2 | 11.8 | 49  | 731 | 941 outcropping | 0 marginal   | uncertain valley-side Snowdon ignimbrite,    | 40  | 56 Beudy Ma     | 439 | 111 | 33.62962 | 58.4 | 0.946987 | 1.81E+08 | 965.7564 | 479 | 2.818544 | 2.79588  | 2.568202 | 2.642465 | 2.75263  | 2.045323 | 2.680336 | 0        |   |
| 284 | 2628 | 3553 | 588 | 610 | 640 | 916 | 921 | 885 | 502 | 245  | 276 | 52.5 | 8.1  | 12  | 694 | 926 outcropping | 0 marginal   | uncertain valley-side Snowdon ignimbrite,    | 67  | 60 Crib Coch I  | 297 | 52  | 30.60996 | 44.4 | 0.488048 | 3.65E+07 | 331.7993 | 328 | 2.700704 | 2.389166 | 2.440909 | 2.472756 | 2.520675 | 1.716003 | 2.515874 | 0        |   |
| 290 | 3267 | 2344 | 452 | 480 | 521 | 615 | 684 | 615 | 310 | 438  | 108 | 51.3 | 12.5 | 43  | 330 | 407 drift       | 0 marginal   | uncertain valley-side Black Mou Old Red St.  | 79  | 83 Ochon, B.    | 163 | 69  | 27.73568 | 38.8 | 1.412903 | 2.21E+07 | 280.7638 | 163 | 2.491362 | 2.641474 | 2.025308 | 2.212188 | 2.448341 | 1.838849 | 2.212188 | 0        |   |
| 291 | 3239 | 2344 | 503 | 515 | 563 | 642 | 671 | 642 | 382 | 504  | 80  | 43.2 | 7    | 55  | 281 | 333 drift       | 0 marginal   | none valley-side Black Mou Old Red St.       | 60  | 67 Eflengyl S., | 139 | 60  | 19.99515 | 36.2 | 1.319372 | 2.68E+07 | 299.1136 | 139 | 2.582063 | 2.70243  | 1.93039  | 2.143015 | 2.475836 | 1.778151 | 2.143015 | 0        |   |
| 292 | 3254 | 2307 | 392 | 400 | 445 | 595 | 640 | 560 | 380 | 458  | 132 | 60.5 | 3.3  | 57  | 364 | 408 major bog   | 0 definite   | LLS glacie valley-side Black Mou Old Red St. | 39  | 35 Maes-y-ffr   | 168 | 53  | 23.85047 | 57.2 | 1.20263  | 2.92E+07 | 308.0724 | 173 | 2.578784 | 2.660666 | 2.120574 | 2.225309 | 2.488653 | 1.724276 | 2.238046 | 0        |   |
| 293 | 3238 | 2321 | 430 | 440 | 488 | 602 | 680 | 598 | 669 | 1494 | 140 | 65   | 1.5  | 88  | 328 | 392 major bog   | 0 poor       | LLS glacie valley-side Black Mou Old Red St. | 66  | 60 Elogog, Tai  | 168 | 58  | 14.09668 | 63.6 | 2.233184 | 1.68E+08 | 351.6903 | 172 | 2.825426 | 3.174351 | 2.146128 | 2.225009 | 2.741695 | 1.763428 | 2.235528 | 0        |   |
| 294 | 3013 | 2242 | 339 | 390 | 451 | 525 | 590 | 515 | 570 | 575  | 110 | 41.4 | 11.3 | 110 | 403 | 521 drift       | 0 marginal   | uncertain valley-side En. & C. B Old Red St. | 307 | 300 Dyffant LJ  | 176 | 112 | 17.15922 | 30.1 | 1.008772 | 5.77E+07 | 388.3834 | 186 | 2.756875 | 2.759668 | 2.041393 | 2.245613 | 2.587018 | 2.048218 | 2.269613 | 0        |   |
| 295 | 2843 | 2182 | 273 | 300 | 343 | 463 | 635 | 448 | 345 | 760  | 130 | 52.5 | 3.7  | 33  | 379 | 513 drift       | 0 marginal   | uncertain valley-side Wn. Breco Old Red St.  | 71  | 73 S. Tawe F.   | 175 | 70  | 26.89624 | 48.8 | 2.202699 | 4.59E+07 | 358.006  | 180 | 2.537819 | 2.880814 | 2.113943 | 2.243038 | 2.65389  | 1.845098 | 2.278754 | 0        |   |
| 296 | 2837 | 2191 | 420 | 490 | 525 | 664 | 667 | 616 | 450 | 860  | 140 | 50.2 | 5.2  | 29  | 416 | 539 drift       | 0 marginal   | uncertain valley-side Wn. Breco Old Red St.  | 69  | 74 C. Tawe F.   | 196 | 105 | 23.5358  | 46   | 1.911111 | 7.59E+07 | 423.3072 | 244 | 2.653213 | 2.934469 | 2.146128 | 2.282256 | 2.626956 | 2.021189 | 2.38739  | 0        |   |
| 297 | 2834 | 2203 | 515 | 600 | 642 | 761 | 761 | 740 | 485 | 1685 | 140 | 66.6 | 2.4  | 29  | 271 | 518 major bog   | 0 marginal   | LLS glacie valley-side Wn. Breco Old Red St. | 88  | 84 N. Tawe F.   | 225 | 127 | 24.88741 | 64.2 | 3.474227 | 1.84E+08 | 568.6452 | 246 | 2.685742 | 3.2206   | 2.146128 | 2.352183 | 2.754841 | 2.103804 | 2.309035 | 0        |   |
| 298 | 2801 | 2217 | 500 | 502 | 576 | 700 | 701 | 668 | 705 | 890  | 175 | 78.7 | 0    | 110 | 336 | 492 major mori  | 0 classic    | LLS glacie valley-hear Wn. Breco Old Red St. | 29  | 6 L. Y Fan F.   | 168 | 76  | 13.40348 | 78.7 | 1.262411 | 1.05E+08 | 472.385  | 200 | 2.848189 | 2.94939  | 2.243038 | 2.225309 | 2.674296 | 1.880814 | 2.30103  | 0        |   |